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COMPARISON OF PERIODIC

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ROCKET DATA

G.D. Nastrom and A.D. Belmont

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8100 South 34th. Avenue

Minneapolis, Minnesota 55440



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16. Abstract The purpose of this study is to compare the temporal variations in stratospheric winds and temperatures with the geomagnetic field elements. From a periodic analysis of the geomagnetic field elements, based on data from 1960-1972, the amplitude and phase of the quasibiennial, annual, and semiannual waves are given for stations from 1°S to 89°N. These results are then compared with corresponding waves reported in rocketsonde wind and temperature data, 30 to 60 km. The annual waves are found to be coupled as a result of the annual variation in the dynamo effect of the wind in the lower ionosphere. The semiannual waves are also found to be coupled and three possible causes for the extra-tropical stratospheric semiannual wind wave are discussed. Time variance spectra for the interval from 4 days to 44 days in both zonal winds and horizontal geomagnetic field intensity are compared for years when major midwinter warmings occur and years when only minor warmings occur. The noted differences are suggested to arise from upward propagating planetary waves which are absorbed or refracted in varying amounts depending on the prevailing circulation. Lastly, a superposed epoch study reveals a statistically significant correlation between stratospheric temperature and k_p fifteen hours earlier. The possible reason for this peak is discussed, but a similar relationship with respect to the solar sector structure could not be found.			
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I. INTRODUCTION

Coupling between the atmospheric circulation and the earth's magnetic field is strongly suggested by the evidence presented in the literature. This evidence covers a wide spectrum of space and time scales, is usually in the form of correlations, and has been given for all levels of the atmosphere. For example, Flohn (1952) showed that the meteorological equator (the ITCZ) is more nearly parallel to the geomagnetic than to the geographic equator and that the polar vortex at 200 mb is more nearly centered on the geomagnetic than on the geographic pole. King (1974) has shown that the isolines of total ozone are similar to the isolines of magnetic field strength; and Belmont, et al. (1974b), showed that the contours of the amplitude of the semiannual wave in zonal wind at 50 km are more nearly congruent with the geomagnetic, rather than geographic, coordinate system. The mechanisms which give rise to these correlations are not yet fully understood.

It should be determined whether the atmosphere or the geomagnetic field, or neither, is the independent variable responsible for correlations such as the above. If the geomagnetic field is the independent variable for a given relationship, then meteorologists ought to include that relationship in their studies. For example, a high latitude source of NO produced by cosmic rays, which enter the atmosphere at latitudes determined by their interaction with the geomagnetic field, is now being included in studies of the ozone budget (Crutzen, et al., 1975). On the other hand, if the atmosphere is found to be the independent variable for a given relationship, then such knowledge may be useful to space scientists, but meteorologists need not consider that geomagnetic relationship in studies of the atmospheric circulation.

The purpose of the present study is to find relationships between stratospheric parameters, 30-60 km, and geomagnetic field parameters. The mid- and upper-stratosphere may respond dramatically to geophysical events (e.g., Zadvernyuk, 1973), and by studying relationships at high altitudes

it may be possible to more readily identify some coupling mechanisms between the atmosphere and geomagnetic field. Hopefully, this could help explain which are the independent variables for some of the known relationships between geomagnetic and meteorological data. The method used here will be to compare temporal variations of wind and temperature at rocketsonde stations (MRN data) with time variations of the vertical (Z) and horizontal (H) geomagnetic field intensity at a nearby geomagnetic observatory.

The frequency range of time variations which can be studied is limited only by the time distribution of MRN data, as the geomagnetic data are taken hourly in a (usually) continuous sample. The MRN data are sufficiently plentiful to define variations longer than a month, so a major portion of the study deals with periodic analysis, the quasi-biennial oscillation, and the first three harmonics of the annual wave. On time scales of a season or less, midwinter sudden stratospheric warmings are the most spectacular events. While the MRN data are too sparse to perform case studies of individual warmings, it is possible to stratify all years according to whether or not a major warming occurred. This procedure has been used to study differences of the variance spectra in MRN and geomagnetic data during years when major warmings occurred compared with the other years. Finally, results are given for a superposed epoch study of the changes in stratospheric temperature over a few hours time at Fort Churchill using solar sector boundary crossing dates as the key events.

II. DATA

Meteorological rocket (MRN) data 1960-72 were obtained from the World Data Center, Asheville. Station locations and the nearest geomagnetic observatories are given in Table 1. Further details concerning the MRN data and results of periodic analysis of wind and temperature have been given in Belmont, et al. (1974a), and Nastrom and Belmont (1975).

Daily mean values of the geomagnetic field elements for years corresponding to the MRN data were obtained from the World Data Center, Boulder.

Observatories used in this study are listed in Table 2. Generally, geomagnetic data prior to 1960 were not used in order to make the periods of record of the MRN and geomagnetic data as compatible as possible. Also, the analysis was limited to observatories near a MRN station (see Table 1). At most observatories the field elements given are declination (D), horizontal (H), and vertical (Z) field intensity. At the Canadian stations indicated in Table 2 they were reported as X, Y, and Z; but daily values were converted to D, H, and Z prior to further processing. The observatory at San Juan was moved at the end of 1964, and it was necessary to adjust the baseline of 1960-1964 data to be consistent with 1965-1972 data. Also, the observatory at Honolulu was moved during 1960, so data for 1960 were not used there.

III. RESULTS

A. PERIODIC ANALYSIS OF GEOMAGNETIC FIELD ELEMENTS

1. Procedure

Significant peaks at 12 and 6 months are present in the H and Z spectra (Currie, 1966), and resolving them as discrete lines provides the possibility of studying their phases as well as amplitudes. However, secular trends are often very large in the H and Z data, and failure to remove them prior to periodic analysis can lead to inconclusive results (Chapman and Bartels, 1940, Chapter 16; Currie, 1966). Inspection of plots of our time series for 1960-1972 (not shown here; see Chapman and Bartels, 1940, p. 132) indicated that a parabola can be used to effectively remove the secular trend. This technique is more desirable than other filters because no data are lost at each end of the time series. Although a parabolic trend line does interact with the approximately 11 year cycle found in H and Z, numerical tests made using synthetic time series show that the error in amplitude and phase of the 11 year cycle, and shorter periods, is less than 4% after removing a parabolic trend from a time series 13 years long.

Time series of mean daily H and Z values are characterized by a

relatively steady background which may be dramatically interrupted during a geomagnetic disturbance. The effect of geomagnetic storms, which last for only hours or days, could be thought of as a very large amplitude high-frequency variation which occurs more during some years than during others. As the purpose here is to study the month-to-month changes of the background geomagnetic field, it is desirable to remove the aliasing of monthly data caused by the irregular occurrence of geomagnetic storms. One method to achieve this is to use only non-disturbed days when computing monthly means. This method has the drawback that disturbed days can be identified only on a subjective basis. If the disturbed days form only a small part of total daily values in a month, however, then an objective and nearly as effective method is to use monthly medians rather than monthly means. Periodic analyses were made using both monthly median and monthly mean data. The resulting amplitudes and phases differed significantly between the two analyses with the monthly median amplitudes always smaller (e.g., 5.3 versus 8.8 gammas for the amplitude of the annual wave in H at College). Moreover, the corresponding statistical error estimates were always smaller in the case of monthly median data, indicating less interannual variability of the periodic waves when monthly medians are used. Thus, the following analyses are based on time series of monthly median values of H and Z. (Note that monthly mean MRN data were used here, as previously, because the range of those fluctuations is relatively much smaller.)

Monthly data of both the MRN and geomagnetic parameters were analyzed with the joint periodic regression technique of Belmont, et al., 1974a. The technique can be used to analyze a time series of irregularly spaced data points, weighting the months by number of observations, even if zero, and to simultaneously determine an estimate of the statistical error of the amplitude and phase of each frequency included. Further, frequencies analyzed need not be integral divisions of the period of record. Frequencies included in the present analyses are the long-term mean, 11 year cycle (geomagnetic data only), quasi-biennial oscillation (29 months), and the first four harmonics of the annual wave.

The statistical errors given in Table 2, which provide confidence estimates for the results, are not the same as RMS deviations. However, they resemble RMS deviations because the coherence of the data from cycle to cycle is the most important consideration in computing them; in fact, regression of the errors in Table 2 (SE) for the annual and semiannual waves with RMS deviations determined from harmonic analysis of yearly data showed that $SE = 0.38$ RMS. The regression coefficient is small for several reasons: because the frequencies included in the SE analysis are not orthogonal over the data, they can interfere with each other to give a better fit (smaller residuals) to the complete time series than can the orthogonal components of harmonic analysis. Also, as SE weights each point of the time series by the number of observations, it allows occasional erratic points based on few observations to be largely disregarded.

2: Annual Wave

Estimates of the amplitude and phase, with errors, of the QBO, annual, and semiannual waves in H and Z are given in Table 2. Results for the annual wave in H and Z are plotted in Figure 1 as functions of geomagnetic latitude. Lines estimating the latitudinal variation in the figure have been fitted by eye. These results are similar to the corresponding values given by Currie (1966), but have much less scatter, particularly at mid-latitudes. As noted by Currie, the scatter of his results may arise from differing periods of record at the various stations he used; for example, the interannual variations of the annual waves in H at Tucson and Sitka (Figure 2) are so large that averaging a given number of arbitrary years will clearly lead to widely varying mean values. In anticipation of the discussion in Section B, a large part of the interannual variations in Figure 2 is due to the well-known solar-cycle influence on E region ionization. Note that both stations are at the right extreme in 1963 and the left extreme in 1969 (1963 was near sun spot minimum and 1969 near sun spot maximum). Returning to Figure 1, the sharp increase of amplitude of the annual waves in H and Z at high latitudes has been

noted by Currie (1966) whose data extended to 80°N (geomagnetic latitude). The decrease of amplitude in H and continued rise in Z as the pole is approached does not seem to have been reported previously, although Langel and Brown (1974) have noted that the largest seasonal variations of ΔZ are near the pole.

The phase of the annual wave in H is fairly uniform at all latitudes, with the annual maximum occurring in June (Figure 1a). The phase of Z , on the other hand, undergoes an abrupt shift of 180° near 65°N . Equatorward of 65°N the average phases of Z and H are quite similar.

3. Semiannual Wave

Amplitudes and phases of the semiannual variations in H and Z are plotted in Figure 3. The present amplitude results are generally smaller than those given by Currie (1966). In this case, the difference may again be due to differing periods of record, or it may be due to our use of monthly median data which reduces the occasionally severe impact of magnetic storms which occur on a predominantly semiannual rhythm. It should be borne in mind that the present results are for the mean semiannual wave over about one sunspot cycle. [Chapman and Bartels (1940) have shown that the semiannual amplitude varies with the sunspot cycle.] From 70° to 80° magnetic latitude, the decline of amplitude in H and the increase in Z are in accord with Currie's (1966) results which extend to Godhavn (80°N). As the pole is approached from 80°N the amplitudes of both H and Z increase, although the large statistical errors associated with the H values make that analysis less reliable.

The phase of the semiannual variation in H is fairly steady up to about 70°N , with a 180° shift near 75°N indicated by all three stations north of 80°N . The phase of Z is less steady, but indicates a systematic shift with latitude such that the phase of the pole and the equator are about 180° different.

4. Quasi-Biennial Oscillation (QBO)

Before discussing the periodic analysis results for the QBO in H and Z, it should be pointed out that there have been several conflicting reports regarding the existence of a quasi-biennial "line" in the geomagnetic spectrum. Hope (1963) reported that the QBO in K_p had been isolated, but Currie (1966) could not find it in the spectra of H or Z and suggested that these results were based on faulty numerical filtering procedures. Fraser-Smith (1972) presented the spectrum of the A_p index and concluded that no QBO exists, but Currie (1973) has analyzed H and Z data from 49 observatories and now concludes that there is a line near 2.15 years.

Nearly all periodic waves in geophysical data show variations from cycle to cycle, but usually the amplitude and phase converge on mean values if enough cycles are averaged. Statistical tests can be used to determine if enough cycles of a periodic wave have been used to estimate the mean wave with confidence (Chapman and Bartels, 1940). As the quasi-biennial oscillation is not truly periodic, but has variable amplitude, phase, and period from cycle to cycle (e.g., see Figure 4), there is no assurance that a mean wave can be rigorously defined in a usual statistical sense. Thus, the mean QBO can only be defined for the years of record analyzed by each writer, with the understanding that the QBO for different years of record will probably not have the same amplitude or phase. For this reason, the latitudinal variation of the QBO values in Table 2 is erratic and inconclusive unless stations with the most complete and most nearly identical years of record are considered. Therefore, only those stations with over 110 months of data have been used in order to obtain the most reliable estimates of the latitudinal variation of the QBO. (The average period of the QBO during these years was 29 months.) In Figure 5, the average QBO amplitude is near 2 γ for both elements, although Resolute (83°N) indicates an increase of H's amplitude as the pole is approached. The individual phase dates, relative to 1 January 1960, have fairly uniform latitudinal variation except for Z at College

and H at Resolute.

B. COMPARISON OF GEOMAGNETIC AND METEOROLOGICAL PERIODIC WAVES

The present objective is to determine possible relations of geomagnetic to meteorological variables by comparing periodic properties of geomagnetic and MRN data. Identifying those periodic frequencies which show a close relationship can allow effort to be focused on them, with the remainder of the variance discarded as being unrelated and therefore of no immediate interest.

1. Procedure

Waves of the same period whose relative phase lags show broad patterns of spatial continuity are sometimes found to be related, and charts of the relative phase lags between MRN and geomagnetic periodic variations will be presented below. However, as all periodic components of the two data sets may have large year -to-year variability in amplitude and phase (for example, the annual wave in H in Figure 2), it is desirable to first examine the year-to-year relationship of each frequency to determine if "average" phase lag values are representative. The coherence square (COH2) statistic of cross spectral analysis provides an objective measure of how uniformly the amplitudes and phases within each frequency band vary with time at a given location, and has been used here to decide which frequencies of MRN and geomagnetic data are synchronous. Cross-spectral analyses of horizontal and vertical components of the geomagnetic field versus the temperature and wind at nearby rocket stations were made using monthly data at five stations with the most complete periods of record (Table 3), yet well distributed in latitude, and with a maximum lag of twelve months. Prior statistical significance of each COH2 value was tested by the method of Julian (1975). No values in the frequency band centered at 24 months (near the QBO frequency), nor more values than expected by chance for frequencies higher than $2\pi/6$ months, passed the 5% confidence level. COH2 values for frequency bands centered

at biennial, annual, semiannual, and terannual periods are given in Table 3; those which exceed the 5%, 1%, and 0.1% confidence levels are marked.

There is apparently closest coupling for the annual and semiannual variations of zonal wind with H at mid-latitudes and with Z at lower latitudes. The annual temperature variation, especially at 30 km, also is significantly coupled with geomagnetic variations. Semiannual variations of temperature and all terannual variations exceed the 5% confidence limit no more often than expected by chance. These results indicate that only the annual, and some semiannual, variations in MRN and geomagnetic data exhibit significantly synchronous year to year changes; therefore, only those waves will be considered further.

There are two techniques for determining relative phase lags: first, relative phase lag can be found during cross-spectral analysis. Second, the phases determined by periodic analyses can be subtracted. The latter method has the advantage that a measure of confidence can be derived by combining the statistical phase errors determined during periodic analysis. This was done by the root-sum-square technique. It was found that all phase lags associated with a COH2 in Table 3 which exceeded the 1% confidence limit were within the limits of statistical error of the phase lags determined by subtracting periodic analysis results. Therefore, values presented below are based on periodic analysis results.

2. Relative Phase Lags

Relative phase lags for the annual and semiannual waves between MRN and geomagnetic data are presented in Figure 6 as functions of height and latitude. For each station pair listed in Table 1a, the relative phase lag of each frequency for each parameter was determined by subtracting the phase of the geomagnetic wave from the meteorological wave, at 4 km height intervals from 28-64 km. The resulting phase lag values were plotted at the geographic latitude of the MRN station. Contours were drawn for phase lags = ± 30 , ± 150 degrees to indicate areas of nearly in

or nearly out of phase. The relative uncertainty of each value, estimated by the root-sum-square of the individual phase errors, and the spatial patterns of phase given in Figures 1 and 3, and in Belmont, et al. (1974a), and Nastrom and Belmont (1975), were taken into account while drawing the contours.

In Figure 6, U-H are out of phase throughout the mid-latitudes for both the annual and semiannual waves. U-Z are out of phase from about 10°N - 40°N for the annual wave, with small in phase areas at high latitudes. The phase lags of U-Z for the semiannual wave are near 180° in the upper tropical stratosphere and nearly in phase at high latitudes. The annual waves in T-H are out of phase above 55 km near 20°N , and in phase north of a line from 28 km, 10°N to 64 km, 60°N . The annual waves in T-Z appear out of phase in the upper low-latitude stratosphere and at highest latitudes, and are in phase near 30 - 50°N . Clearly, the phase lags presented in Figure 6 have broad spatial continuity. Together with the large COH₂ values these results suggest that physical coupling between the MRN and geomagnetic periodic variations may exist. Possible mechanisms which could produce coupling will be discussed next.

3. Discussion

a. Annual Wave

It must be noted here that the annual variation in geomagnetic data is not yet fully understood, although several writers have discussed it. Vestine (1954) suggested it could be a seasonal effect induced by air motions in the ionosphere. Currie (1966) concurred with Vestine and offered qualitative arguments from the scanty data then available, and later (Currie, 1974) strengthened the theory by arguing the annual wave could not arise from modulation of the S_q current system but must be a DC effect.

(1) Suggested Mechanism: Due to the differing ion and electron Hall conductivities, zonal wind in the lower ionosphere produces a ring current along the wind which induces a magnetic field in the meridional

plane. This induced magnetic field affects the geomagnetic field in proportion to the wind speed, and the effect decreases with distance. The maximum effect on the N-S component of the geomagnetic field will be directly below and above the wind jet where the induced field is coincident with the geomagnetic H field. Similarly, there is a maximum effect on the Z component to the north and south of the zonal wind jet with a minimum directly below and above it.

At high latitudes (Fig. 1), the maximum amplitude of the annual wave in both H and Z occurs. It apparently has not yet been explained in the literature. The cause of the high latitude maximum could be the annual variation in ionization density, which is a function of solar elevation angle, and winds in the lower ionosphere or of magnetospheric origin; but there is insufficient data to verify either hypothesis at this time.

At mid- and low-latitudes, however, sufficient data are now available to crudely estimate the magnitude of the annual effect of ionospheric winds on geomagnetism and thereby perhaps bring future research efforts on this issue into focus. Here the annual variation in wind is the major factor as there is only a small seasonal change in electron density. The annual variation in zonal winds in the lower ionosphere has maximum amplitudes of about 30 m/s from 20-50° latitude near 110 km with phase dates near mid-May (Groves, 1972). Ionized gas is dragged eastward during the half year centered about May, and westward during the half year centered about November, producing an annual variation in the geomagnetic field intensity. To estimate the magnitude of this effect, the current sheet approximation is applied using a width of 500 km (after Bates, 1975), depth of 15 km, uniform charge density of $5 \times 10^4 \text{ cm}^{-3}$, at an altitude of 110 km. A wind variation of (30 m/s) $\cos(\omega t + \phi)$ yields a field variation of (3.28 gammas) $\cos(\omega t + \phi)$. Of course this estimate could easily be changed by a factor of two or more, but the amplitude is certainly of the proper order of magnitude for the mid-latitude annual wave in geomagnetism (Fig. 1). Also, the charge density

varies, particularly with solar zenith angle; this could account winds (mid-May) and the geomagnetic variation (mid-June). Since the annual amplitude in the zonal wind above 100 km has a maximum near 50°N it should create a maximum in the annual amplitude of H near 50°N and a minimum in Z near 50°N, as found in Figure 1.

Conventional heat sources (e.g., radiative heating) are adequate to account for the annual wind waves in the stratosphere (Leovy, 1964) and lower thermosphere (Volland and Mayr, 1972). The MRN individual data has high coherence with the geomagnetic data at the annual frequency because the variations in annual wave between the thermospheric and stratospheric wind are apparently also coherent. (The fact that the annual wave in the stratosphere is out of phase with that in the thermosphere has no bearing on their coherence.) The point here is that a seemingly intriguing relation between two parameters may arise from a mutual association with a third parameter through normally accepted processes; in this case the third parameter is the annual wave in thermospheric circulation. Hence, the present results do not suggest any geomagnetic influence on the atmospheric circulation.

These results should be useful to those trying to understand apparent correlations between atmospheric and geomagnetic processes, and to those concerned with the description of the earth's magnetic field and its variations. The annual dynamo concept presented above could be incorporated into models of the geomagnetic field and thereby help overcome the problems of interpretation discussed by Alldredge and Stearns (1974).

As the above calculation, based on constant ion density, does not pertain to the large, high latitude annual waves in geomagnetism, it is not inconsistent that relatively low COH2 values for the annual frequency are found in Table 3 at Greely and Churchill. Values of COH2 in Table 3 at mid- and low-latitude stations are less than 1.0 for reasons

besides instrument error and incomplete sampling. Solar cycle influence on charge density in the ionosphere, causing a solar cycle in the annual wave in geomagnetism but not in that in stratospheric wind, may be the most important additional reason. However, upward propagating planetary and gravity waves, which may affect the stratosphere and ionosphere much differently, could also be important. Although recent theories suggest that planetary waves will be absorbed, reflected, refracted and radiatively damped in the stratosphere and mesosphere, there is a large body of evidence which suggests they do exist in the lower thermosphere (e.g., Lysenko, et al., 1974; Deland and Friedman, 1972; Graznik, et al., 1975). The possible role of gravity waves in the upper atmosphere is also poorly understood (Muller and Kingsley, 1974). It seems unlikely that these uncertainties will be cleared up until detailed wind measurements from the surface to the lower thermosphere are studied. A preliminary effort has been made by Manson, et al. (1975), but conclusive results are not yet available.

(2) Applicability to Correlation Studies: During an early phase of the present study the linear correlation coefficients between the monthly means of MRN and geomagnetic data were computed. Those results, given in Table 4a, have a high level of statistical significance. It is now realized that the correlation coefficients are large because the annual waves in MRN and geomagnetic data are coupled and, except at low latitudes, the annual wave is generally larger than any other periodic component in the MRN data. Thus, one would expect the linear correlation between zonal winds and geomagnetic data to decrease significantly if the annual waves were removed from both data sets. To test this hypothesis, the linear correlation coefficients were recomputed between the monthly residuals after the annual waves had been subtracted. The results of this test, given in Table 4b, show that in nearly all cases the correlation ceases to be significant when the annual wave is removed. The correlation remains significant at Hawaii because the semiannual wave in zonal wind is nearly as large as the annual.

Application of this point to other reported correlations may help explain them. For example, King (1975) has reported that the longitudinal variations at 60°N of the average 500 mb height for January and the geomagnetic intensity shifted 25° in longitude have a correlation coefficient of -0.963. Longitudinal variations in the circulation of the mid-stratosphere reveal a standing wave up to at least 10 mb; in the meridional component the predominant standing wavenumber is two (van Loon, et al., 1972; Figure 72). Lysenko, et al. (1972), and Glass, et al. (1975), have offered evidence that standing waves also exist in the circulation of the lower thermosphere. If the predominant wavenumber of wind speed, ion density, or a combination of them in the lower thermosphere in January is two, then the resulting current will induce a wavenumber two pattern in the longitudinal variations of the geomagnetic field intensity. The high correlation found by King may therefore reflect a very mundane relationship, as long proposed by Wulf (1945), rather than any solar-terrestrial effect. A similar principle could apply regarding the relationship between spatial variations of tropospheric temperature, humidity, and surface pressure and the geomagnetic field intensity reported by King (1974).

b. Semiannual Wave

The results in Table 3 and Figure 6 suggest that the semiannual waves in MRN zonal wind and geomagnetism are also closely coupled. For the zonal wind in Table 3, significant COH2 values are found for the semiannual variation at nearly the same station-levels as for the annual variation. A dynamo mechanism might be suggested, as Groves (1972) shows that there are large semiannual wind variations near 115 km. However, Volland and Mayr (1972) found that most of the latitudinally varying part of the semiannual wind wave in the lower thermosphere is due to corpuscular heating. They suggest (Mayr and Volland, 1971) that this heating is related to the semiannual occurrence of magnetic storms, which Chapman and Bartels (1940) have argued is due to earth-sun geometry and thus is independent of meteorological influence. There-

fore, the close coupling of the semiannual waves seen in Table 3 and Figure 6 may be explained independently of the dynamo mechanism of the annual wave.

More insight regarding the cause for this coupling of the semiannual waves in geomagnetic and MRN data might be possible if the cause of the semiannual wave in zonal wind were known. Possible causes for the tropical semiannual wind wave have been discussed by Dickinson (1975), but the extra tropical semiannual wave has not yet been explained. As processes which show more symmetry in one coordinate system than another may be driven by mechanisms peculiar to that coordinate system, tests of the relative symmetry of the semiannual wind wave in geomagnetic and geographic coordinates were made. These tests, described below, were generally inconclusive. Finally, three possible causes of the extratropical semiannual wind wave are discussed. None can yet be accepted, and it is suggested that more research is needed before a conclusion can be reached.

(1) Further Tests for Coupling with the Geomagnetic Field: In order to test the relative symmetry of the semiannual wind wave in geomagnetic compared with geographic coordinates the relative phase lags of Figure 6b and 6d have been plotted in geomagnetic coordinates in Figure 7. In either case, the change of coordinates makes little difference, although for U-Z the contours become smoother in geomagnetic coordinates.

Belmont, et al. (1974b), compared the symmetry of the amplitude of the semiannual wind wave at 50 km on maps in the geographic and geomagnetic coordinate systems and found the symmetry slightly greater in geomagnetic coordinates. Even greater symmetry may be found by plotting the amplitude at each station at that level where the closest relationship is found. The height of the level at each station was selected as that height where the magnitude of the product of the two semiannual waves' amplitudes and the cosine of their phase lag ($a_2 \cdot b_2 \cdot \cos\Delta\phi$) is maximum.

Note that this parameter, which is an approximation of the co-spectrum in the case of large COH2, will be relatively small if either amplitude is small or if the phases are near quadrature. The height of this surface is shown by the dotted line in Figure 6b. Contours of the amplitude of the semiannual wind wave at the heights thus selected are shown in Figure 8, and appear to show little, if any, enhanced symmetry in either coordinate system compared with the results of Belmont, et al. (1974b). Clearly, these tests for increased symmetry do not suggest preference for either coordinate system.

(2) Possible Mechanisms: Three hypotheses can be advanced to account for the extratropical stratospheric semiannual wave in zonal wind. Before discussing them, however, it should be pointed out that Gregory, et al. (1975b), have noted that the phases of the annual wind waves in the stratosphere and upper mesosphere are reversed. Cole and Kantor (1974) have noted a similar relationship with regard to the annual waves in temperature at stratospheric and mesospheric levels. Both papers suggest that the semiannual waves in the lower mesosphere at extratropical latitudes result from the overlapping of the annual waves. This descriptive account of the lower mesospheric semiannual wave is useful, but does not by itself explain the semiannual wave. For example, early descriptions of the tropical semiannual wind wave in the upper stratosphere viewed it as the result of alternating intrusions of winter hemisphere westerlies into the summer hemisphere (Webb, 1966). While that does occur, it does not explain the tropical semiannual wave, and efforts to do so have invoked a wide variety of mechanisms, e.g., ozone heating, the diurnal tide, planetary waves, Kelvin waves, and semidiurnal tides. Hopefully, the discussion below will help stimulate other research efforts to explain the extratropical semiannual wind wave.

The first forcing mechanism for the extratropical semiannual wind wave to be considered here is upward propagating planetary waves. If these waves interact with the background flow on a semiannual basis they could induce a semiannual component in the background wind speed.

In an effort to determine if the amount of absorption of planetary waves varies with season, the variance spectrum of filtered time series of zonal and meridional winds have been determined at eight MRN stations on a seasonal basis. Details of procedure and complete results are given in the Appendix. As noted in the Appendix, maximum power usually occurs between $2\pi/10$ and $2\pi/20$ days, so for brevity only the results at $2\pi/11$ days will be presented below. However, the graphs for the total variance and for the results at $2\pi/6.3$ days (not shown) are similar.

Planetary waves propagating vertically in a hydrostatic atmosphere with no dissipation increase their spectral density (i.e., power) exponentially with height. Further, if attenuation of the waves occurs, the slope of the power will be proportional to the amount of attenuation. Spectral density for the frequency band centered at $2\pi/11$ days is presented as a function of height in Figure 9 for six MRN stations. Although the power at a given level changes with season, sometimes dramatically, the slope of the curves does not change much with season except at Kennedy. There is a large difference at Kennedy between the slope during spring and autumn from that during the solstitial seasons. During winter unexplained absorption occurs above 45 km at both Kennedy and Pt. Mugu. These results support the hypothesis of Belmont, et al. (1974b) that planetary wave absorption is responsible for the secondary amplitude maximum of the semiannual wave found near 30°N . They do not, however, suggest that the semiannual wave at other latitudes arises directly from seasonal absorption of planetary waves.

The second proposed mechanism is influence on the ozone field by particle precipitation. This mechanism was proposed by Belmont, et al. (1974b), but cannot yet be directly tested due to a dearth of high level ozone data. However, it should be noted that Heath (1974) has found evidence for a non-photochemical source of high latitude ozone creation which he attributes to incident charged particles; and recent modeling efforts by Crutzen, et al. (1975), have shown that incident charged particles can dramatically influence the ozone field. Also, Golyshev,

et al. (1974), found that the amplitude of the semiannual wind wave near the stratopause exhibits a solar cycle modulation. To illustrate this, yearly values of several solar and geophysical parameters are presented in Table 5. A station-year is not included in the table unless data for all twelve months are available, and temperature data were thus too irregular to include in the table. Note that the values of the sunspot number and of the semiannual amplitudes have relative maxima in 1969 in all cases. Further, note that the annual wave in zonal wind is a relative minimum in 1969 at all stations except at Barking Sands. While this table suffers from the short period of record available, it does support and extend the results of Golyshev, et al. (1974).

Solar cycle modulation of the periodic variations in stratospheric zonal wind, as seen in Table 5, is consistent with the hypothesis that particle precipitation during magnetic storms influences the ozone and hence thermal and wind fields. If stratospheric semiannual wind variations are related to the occurrence of magnetic storms through the ozone field, then their amplitude should be largest during active sun years (as found in Table 5) as the semiannual component in magnetic storm frequency is largest during active sun years. Solar cycle modulation of the annual wind wave is not easily conjectured, but the well-known solar cycles in total yearly magnetic storms and yearly solar flare occurrence may prove responsible, especially in view of the results of Crutzen, et al. (1975), regarding particle precipitation and ozone concentration.

The third possible mechanism is IR radiation generated in the lower thermosphere during magnetic storms and absorbed by CO_2 at 30-40 km. During magnetic storms the amount of IR radiated by the lower thermosphere is increased by several orders of magnitude, and Gordiyets, et al. (1972), have suggested that it causes heating of CO_2 at 30-40 km and H_2O at 7-12 km. This mechanism has appeal because 30-40 km is the region where maximum amplitudes of the semiannual wave in observed temperatures occur

(Nastrom and Belmont, 1975), and the semiannual component in the occurrence of magnetic storms would produce the proper phase and periodicity. Large amounts of radiative energy are possible for brief periods during severe geomagnetic disturbances, but following Volland and Mayr (1972), the long period form of this heat input (averaged over space and time) should take the same form as the variation in magnetic energy, which is given by

$$U_1^2 \sim \bar{U}^2 [1-0.2 \cos(W_{sa}t)],$$

where \bar{U} is a yearly mean magnetic energy, dependent on solar activity, W_{sa} is the semiannual frequency, and t is time. This implies that the amount of energy deposited by IR radiation is $E_{IR} \sim \bar{E}_{IR} [1-0.2 \cos(W_{sa}t)]$, where \bar{E}_{IR} is a yearly mean value. Alternatively, because layer mean temperature and energy are directly related, $T_{IR} \sim \bar{T}_{IR} [1-0.2 \cos(W_{sa}t)]$. This latter relation says that the IR should contribute five times as much to the mean temperature as it does to the semiannual component of temperature. However, in order to produce a zonal wind oscillation of 20 m/s the latitudinal variation of the corresponding temperature oscillation must be near 5°K (Groves, 1972), which implies a contribution to the average temperature of 25°K . As dynamic models of the stratosphere have encountered no evidence of such a large unconventional heat source (Leovy, 1964), it seems highly unlikely that IR radiation from the lower thermosphere is an important forcing mechanism for the stratospheric semiannual zonal wind wave.

c. QBO

Coupling between the QBO's in MRN and geomagnetic data may exist despite the lack of statistical significance of the COH2 values in Table 3. Even in the tropical stratosphere, where the QBO is the dominant oscillation, the QBO is not regular in amplitude or period from cycle to cycle nor between levels during the same cycle (Wallace, 1973). Thus, the small

COH₂ values may be misleading in this case. Moreover, the QBO in thermospheric zonal winds found by Sprenger, et al. (1975), suggests that the geomagnetic QBO may result from a dynamo mechanism, parallel to the annual wave. Although the present theory explaining the well-known tropical stratospheric QBO appears successful (Dickinson, 1975) it is dependent on waves and processes unique to the tropics and thus cannot be invoked to explain an extra tropical thermospheric QBO. Similarly, any explanation for the thermospheric QBO cannot be based on processes unique to the thermosphere because the large negative correlation between the multi-year variations in Z at Honolulu with the 56 km zonal wind at Barking Sands (Fig. 4) suggests the oscillation is not unique to thermospheric (dynamo) altitudes. Until the altitude and latitude progression of the QBO throughout the upper atmosphere is better known, no conclusion regarding the present results seems warranted.

C. SPECTRAL CHANGES DURING SUDDEN WARMINGS

In the stratosphere, rapid changes in the number, amplitude and phase of planetary waves are the major events during winter. These changes are sometimes associated with "sudden warmings" and it seems of interest to study the changes in MRN and geomagnetic parameters during these disturbances. For this purpose, all years have been categorized as either major sudden warming years (SW) or as other years (MSW). A sudden warming is defined to occur when there is a "reversal of the polar circulation at 10 mb. (30 km) or below". During 1961-72, SW were in 1962-63, 1965-66, 1967-68, 1969-70, and 1970-71, according to a list by Finger.

Several different methods could be used to study the changes of parameters during SW. For example, as planetary wave activity and other events associated with a SW are global in nature (Quiroz, et al., 1975), spatial wavenumber analysis of global data may be used to detect changes in the planetary wave patterns. However, the MRN and geomagnetic

data are not sufficiently distributed geographically to permit detailed spatial analysis. Superposed epoch studies are often useful for single station analysis, but in the case of sudden-warmings it is difficult to meaningfully define a key-date. Indeed, the criteria used for defining the occurrence of a SW are admittedly arbitrary. Thus, the approach used here is to perform power spectrum analysis of single station data and to compare the spectra of SW years with those of MSW years.

The available wind observations (Hook, 1972; Gregory and Manson, 1976) indicate that the circulation of the lower thermosphere is disturbed during a SW. Winds in the ionosphere can act as electric currents and can thereby produce variations in the geomagnetic field. Of course, processes unique to the magnetosphere can also produce variations in the geomagnetic field; but if a geomagnetic spectral feature can be associated with a meteorological process, it may be reasonable to assume that it arises from that meteorological process. Thus, studying spectral changes in the geomagnetic field between SW and MSW years may help better understand the thermosphere. Spectral analysis results for the zonal winds and for the horizontal field intensity are presented first, with a brief discussion of noteworthy features. A comparison of the two sets of results follows and a possible interpretation is suggested.

1. Stratospheric Zonal Winds

Spectra for the zonal winds at 40 km at Fort Greely and White Sands are given in Figures 10-11 for autumn through spring. These stations were chosen because they have the most complete data at high- and mid-latitudes, respectively. In Figures 10-13, K is wavenumber, solid lines are for SW spectra, and dashed lines are for MSW spectra. In autumn and winter there is more energy at Fort Greely (Fig. 10) during SW years if the peaks near $K = 8$ are momentarily disregarded. The high frequency peaks will be discussed later. At White Sands (Fig. 11), however, largest energy occurs during MSW years at $K = 6$ to 9 in autumn and $K = 2$ to 6 in winter. In spring the high frequency energy is significantly

larger at both stations during MSW years. Chi-squared confidence limits have been entered at noteworthy wavenumbers in the figures to indicate the probability that the differences arise from chance. These results for autumn and winter support Matsuno's (1971) suggestions that there is enhanced upward flux of wave energy at high latitudes (e.g., Greely) during SW years, but during MSW years the waves are refracted toward lower latitudes (e.g., WSMR) resulting in more energy there during MSW years.

2. Horizontal Field Intensity

Time spectra of the variations in H at College and Tucson are given in Figures 12 and 13. Spectra for several observatories were computed. As College and Tucson demonstrate the salient features noted and are near the MRN stations used above, only they are presented here. The spectral differences between SW and MSW years noted below are probably due to meteorological influences, and not to solar induced effects. Hauska, et al. (1973), found that over all years 1932-1969 the magnitude of geomagnetic variations in the time range 4 to 40 days varies primarily with the approximately 11-year cycle. During the period 1961-1972, SW years defined above are well distributed over a solar cycle.

In general, at both stations, there is either little difference between SW and MSW years or the spectral values are greater during MSW years. During autumn, the largest differences are at low wavenumbers ($K=1-3$ at CO and $K=1-6$ at TU), while in spring differences are found at intermediate and high wavenumbers ($K=4-5$ at CO and $K=3-11$ at TU). During MSW autumn at College, a significant (1% level) peak is found at $K=9$; less significant peaks at $K=10$ are found there during winter and spring.

3. Discussion

The following chart summarizes which years have the significantly greater spectral values and the wavenumbers at which they occur:

	AUT	WIN	SPR
Greely	SW(3-5)	SW(5)	MSW(3-11)
CO	MSW(1-3)	MSW(1,5-6)	None
WSMR	MSW(6-11)	None	MSW(6-9)
TU	MSW(1-6)	None ..	MSW(3-11)

From the above chart and Figures 10-13 two points should be noted. First, the only time the spectral values are significantly higher during SW years is at Greely during autumn and winter. The first point was noted to be consistent with the theory of wave propagation and refraction. Second, significant spectral peaks near $K=9$ are found only at Greely and College during autumn and winter of MSW years and at Greely during autumn of SW years. The second point may also be explained by planetary waves as will be suggested next.

Planetary waves occur in the troposphere every year. As they propagate upward, they may be refracted toward lower latitudes or they may continue propagating upward, depending on the vertical and horizontal curvature of the flow profile (Simmons, 1974). As waves travel upward they decay; the rate of decay depends on the prevailing circulation. It is now hypothesized that waves of period near 4 to 5 days ($K=8-11$) are upward propagating at high latitudes during all years. If during MSW years they do not suffer severe attenuation then they may continue all the way to the lower ionosphere resulting in spectral peaks near $K=9$ at Greely (40 km winds) and College (lower ionospheric winds). During SW years the prevailing circulation may cause large attenuation or total absorption; thus, during SW years the peak near $K=9$ at Greely in autumn is smaller than during MSW years, and a corresponding peak is not found in winter at Greely nor in autumn or winter at College.

The above arguments, although sketchy and heuristic, are consistent with present knowledge. Lacking from present knowledge, however, is an adequate climatology of circulation differences during SW and MSW years, especially in the upper stratosphere.

D. CORRELATION OF MRN TEMPERATURES WITH K_p AND THE SOLAR SECTOR STRUCTURE

Numerous authors have suggested that the middle atmosphere may be heated following geomagnetic disturbances (e.g., Gordiyets, et al., 1973). A desirable method of testing this hypothesis would be a superposed epoch study using a magnetic storm parameter as the keydate. However, as the MRN data are too scanty to permit that study, the alternative procedure of finding lagged correlations of K_p with respect to the MRN temperature observations was used. The linear correlation coefficients between K_p and the layer mean temperature, 40 - 50 km, at Fort Churchill are given in Figure 14. All temperature soundings, 1960-1972, which had data through the entire layer were used in this study. Values of K_p , obtained from the World Data Center, Boulder, are reported for three-hourly periods; thus, the correlation coefficient was determined at three-hourly intervals as the temperatures were lagged with respect to the K_p values. In Figure 14, negative lag means that the K_p value was measured before the temperature value. The relative maximum correlations are found at lag zero and at lag -15 hours; although both peaks are statistically significant at only the 5% level (if all data are assumed independent), these results are complementary to those found by others.

Ramakrishna and Seshamani (1973) report a statistically significant correlation between the layer mean temperature (from grenade data) at Churchill, 60-89 km, and K_p . The peak correlation occurs when temperature is lagged 15 hours, and the largest correlation coefficients are found when the mean temperature is for the entire layer rather than just the upper portion of the layer. They report this correlation is significant at the 0.1% level. The largest regression coefficients, a measure of the relative magnitude of the effect, are found when only the upper portion of the altitude layer is used; and they suggest this may indicate a larger heating effect at highest altitudes. However, it also may be due to cancellation resulting from opposite effects in different portions of the entire layer. Results given by Zadvernyuk (1973) indicate that the critical layers of the atmosphere may respond

differently to magnetic disturbances; e.g., following a magnetic storm there may be heating at the mesopause but cooling at the stratopause.

Several possible mechanisms could be suggested to account for the correlations discussed above: e.g., corpuscular heating, enhanced IR radiative exchange, etc. However, the correlation could also arise from a meteorological influence on K_p in a manner similar to that suggested by Hines (1973). Therefore, if correlation studies such as the above are to be taken as indicators of a geophysical process important to the lower atmosphere, they must be based on unambiguous parameters so that cause and effect can be clearly discerned.

The previous studies are inconclusive with respect to solar-terrestrial effects. Wilcox (1975) has already related K_p to solar sectors. The real question is whether temperature can be related to solar sector structure. To examine this, a superposed epoch study of the 40 - 50 km layer mean temperatures with solar sector boundary crossings used as key dates is desirable.

However, well defined solar sector boundaries sweep past the Earth at irregular intervals, about every week on the average, and the joint distribution of them with the intermittent MRN observations is not adequate for a superposed epoch study. Thus, it is possible to present in Figure 15 the sign of the temperature change at 40 km between closely spaced consecutive MRN observations at Churchill as a function of time, relative to a solar sector boundary crossing and K_p . The magnitudes of the temperature changes are not shown, but they are random. Solar sector boundary crossing dates were taken from the list in Shapley, et al. (1975). There are 21 temperature rises and 28 temperature falls on the chart, and they seem to be evenly distributed on both sides of the boundary. From this small sample it appears that the temperature trend shows no preference relative to the passage of a solar sector boundary. In summary, either the purported $T-K_p$ relationship is due to some factor other than mutual coupling with solar sector structure, or a much larger sample would be required to establish reliably such a relationship.

IV. SUMMARY

Periodic analysis results of the horizontal and vertical field intensity show that maximum amplitudes of semiannual and annual waves are at high latitudes. It is suggested the high latitude maxima of the annual waves arise from annual waves in ionization density and thermospheric zonal wind speed. At mid- and low-latitudes the annual wave in zonal wind speed in the lower thermosphere (the dynamo region), which is driven by solar heating, accounts for most of the geomagnetic annual wave.

Annual variations in geomagnetic and MRN data are closely coupled. In view of the above cause of the geomagnetic annual wave, coupling between the circulations of the stratosphere and lower thermosphere can explain the geomagnetic-MRN coupling; thus, the present results for the annual wave are not evidence of any geomagnetic influence on the lower atmosphere. Other apparent correlations of the lower atmosphere and geomagnetic field may arise from a similar dynamo action in the thermosphere caused by coupling of the thermosphere with the lower atmosphere.

Semiannual variations in geomagnetic and stratospheric zonal wind data are also closely coupled. As the semiannual wave in thermospheric zonal wind is driven primarily by auroral heating, the cause of this coupling is not clear. It would be helpful if the cause of the stratospheric semiannual wind wave were known, so three possible causes were discussed. Planetary wave absorption seems to be a direct cause only near 30° N, and heating by IR from the lower thermosphere during magnetic storms is energetically unlikely. Possible modulation of the ozone (and hence thermal and wind fields) by particle precipitation during geomagnetic storms has not yet been verified by observations. Amplitudes of the annual and semiannual waves in stratospheric zonal wind may be modulated with the solar cycle as they generally have extreme values concurrent with extreme values of the sunspot number. This result for the annual

wave is believed reasonable as the yearly number of proton solar flares varies with the solar cycle and solar flares can dramatically affect the ozone field. If particle precipitation during geomagnetic storms also influences the ozone field, then this result for the semiannual wave could also be explained as the semiannual variation in geomagnetic activity varies with the solar cycle.

Power spectrum analysis of zonal wind variations shows that at high latitudes there is significantly more wave energy in the upper stratosphere during years when major sudden warmings (SW) occur, but at mid-latitudes largest wave energy is found during years when major sudden warmings do not occur (MSW). This could be explained by wave refraction which occurs in varying degree each year depending on the profile of the background flow; however, a climatology of background flows during SW and MSW years is apparently not available. If geomagnetic variations reflect wind activity in the lower thermosphere then the noted differences between SW and MSW years at mid- and high-latitudes seem consistent with recent theories of planetary wave propagation. The present results suggest that the planetary wave absorption peculiar to SW years occurs in the upper stratosphere, far below the region where direct geomagnetic effects are significant, and thus any direct geomagnetic "trigger" for sudden warmings seems unlikely.

Although the correlation between stratospheric temperature and K_p appears statistically significant and is complimentary to the results of others, cause and effect cannot be discerned. If correlation studies are to be used as evidence of a solar-terrestrial effect, they must be based on parameters of strictly solar origin such as the solar sector structure.

APPENDIX

Organized wave activity in the upper stratosphere has been studied with MRN data by several writers, most recently by Hirota (1975). The latter used only those MRN stations which had at least 30 observations during a given season, subjectively interpolated the data to daily values by analyzing height-time sections for each station, and computed the frequency content of the interpolated data by power spectral analysis. Hirota's method is very effective for analyzing a single season's data; however, the present objective is to prepare a climatology of the power spectra of MRN data, and a less restrictive, objective approach is desirable. Rocket data have historically been taken on an irregular often sporadic basis, and there are instances of many observations at a given station over a time span of a few weeks with relatively sparse data before and after that period. A climatological analysis method should take advantage of those intermittent periods of dense data. The lag-weighted autocovariance function method described below is suited for this purpose, and has been used to estimate seasonal power spectra of MRN wind components, 30-60 km. This method was also used to analyze the power spectra of geomagnetic variations reported in Section III-C of the text. Although a complete description of this method can be found in Dartt and Hovland (1974), a basic outline of it and the variations used here will be given.

A. BASIC DATA HANDLING AND TECHNIQUE

At each 2 km level, 30 - 62 km, multiple rocket ascents over a two-day period were averaged together and counted as one datum in the time series. Due to the poor distribution of MRN observations, many data points represent only one observation and many are missing; but a surprisingly large number (for example, 20% at Churchill) of data points do represent multiple observations. Interpolation was not used for missing data. The time series thus obtained at each station and level were then high-pass filtered by convolution with a discrete, symmetric series of Gaussian weights. To account for missing

observations, the weights under the filter were normalized at each data point such that their sum was always equal to 1.0. The ideal frequency response of this filter is shown in Figure A-1; however, due to missing observations, the actual frequency response is slightly less sharp than shown in the figure.

If an observation is far removed in time from other elements of the time series the filtering process will be ineffective as the datum is then filtered, essentially, with only itself. To preclude this, it was required that there be at least five other data points under the filter (out of a possible 30) and that the sum of the weights before normalization be at least 0.25. These latter conditions resulted in discarding about 10% of the data.

Autovariances up to lag 11 were computed for each individual season, 1961-1972. Three-month seasons were used at all stations with winter defined as December through February. The autovariances and the number of data pairs at each lag and each season were then stored for future use. Seasonal autovariances for all years of record were computed by combining individual seasonal values according to:

$$\overline{R(\tau)} = \frac{\sum R(\tau)N(\tau)}{\sum N(\tau)}$$

where $R(\tau)$ is the autovariance and $N(\tau)$ is the number of data pairs available at lag τ for a given season. With this procedure seasonal autovariances can be computed for the entire period of record, or for just selected years (e.g., years of major sudden warmings). Autovariances thus obtained were Hanned; estimates of the power spectra were obtained by taking the cosine transform of the Hanned autovariances. Finally, the computed values, V , were normalized:

$$\hat{S} = V [2 \text{ MAX } \Delta T / 2\pi]$$

where, in this case, MAX is 11 and ΔT is 2 days; and \hat{S} is the normalized value.

The percentage relative error of each spectral estimate was computed by determining the variance of Hanned spectral estimates according to the formulation of Eddy (1968). The effective number of degrees of freedom required for that calculation were determined with the method of Mitchell (1963). It must be noted that these errors reflect how well each spectrum conforms to a particular statistical model and are only as realistic as that model. Further, they do not account for suspected error sources such as aliasing. Aliasing, or spectral folding, results from sampling at a frequency lower than twice that of the natural variability; this problem is discussed in detail by Lumley and Panofsky (1964). If the true spectrum is a "red-noise" spectrum, as frequency decreases energy increases, then aliasing will tend to make the estimated spectrum flat, i.e., with equal energy at all frequencies. As discussed below, this problem may be more serious for the meridional wind than for the zonal wind.

B. TABLES OF SPECTRAL ESTIMATES

Eight MRN stations have adequate data to provide meaningful estimates of the variance spectrum for the wind components. Temperature observations are less plentiful than wind observations and did not provide useful results. Tables associated with this appendix give climatological spectral estimates for the wind components for each season. The values in the tables have been smoothed with height by a three point binomial filter. As the effect of missing observations on the frequency response of the high pass filter is difficult to estimate, no attempt to restore the spectra has been made; but it appears that the energy in the first frequency (centered at $2\pi/44$ days) is reduced more than the 55% predicted by the theoretical frequency response of the filter.

The relative reliability of the spectrum at each height is indicated by the number of lagged data pairs and the percentage relative error

estimates. The distribution of MRN observations is such that $N(1) \approx \dots$, $N(11) \approx 1/2 N(0)$, so only $N(1)$ is given in the interest of brevity. Also, the percentage relative errors are nearly linear with wavenumber; thus, errors at intermediate frequencies can be estimated by linear interpolation of the values given for $K=1$ and $K=11$.

In spectral analysis there is always a trade-off between resolution and reliability of the spectrum. By averaging adjacent bands a more reliable spectrum may be obtained, but a corresponding loss of resolution results. The most reliable parameter is thus the total variance of the filtered data, which is included in the tables. From experience, the best indicator of the reliability of the total variance is the number of lagged data pairs, and when $N(1)$ is less than about 60 the variance should be disregarded.

In the tables, "VAR" is the total variance of the filtered data, "N" is the number of data pairs at lag one, and "P.R.E." is the percent relative error for bands one and eleven.

C. DISCUSSION

During the Northern Hemisphere summer, the power of zonal wind spectra at a given level generally decreases with increasing K at mid-latitude stations (Kennedy through Wallops) and at Ascension. Zonal wind spectra at Greely, Churchill, and Barking Sands, and meridional wind spectra at all stations, are generally very flat at a given height but have large gradients with height. Although this strongly suggests that aliasing may be a serious problem at the latter stations, Dartt and Hovland (1974) report that summer spectra in the lower stratosphere (30 mb), determined from relatively complete time series of twice daily radiosonde data, are also very flat, especially for the meridional wind. It is therefore likely that aliasing by periods longer than diurnal is not serious in the present results. Note that it is impossible to comment on possible aliasing by periods shorter than diurnal.

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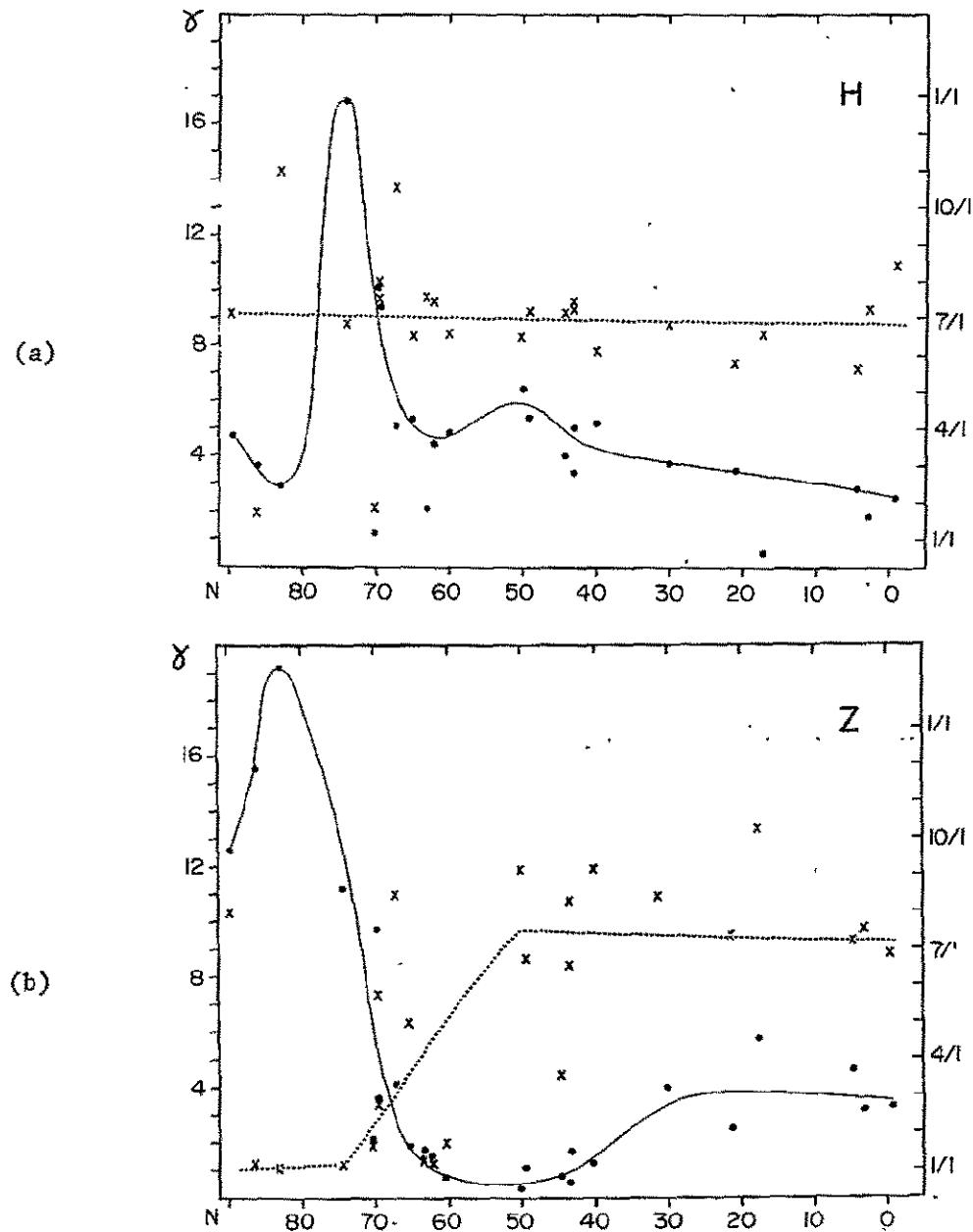


Figure 1. Annual wave in (a) horizontal, (b) vertical field intensity. Dots are amplitudes and crosses are phases of stations in Table 2, plotted at geomagnetic latitude. Lines estimating the latitudinal variation of amplitude (solid) and phase (dotted) have been fitted by eye.

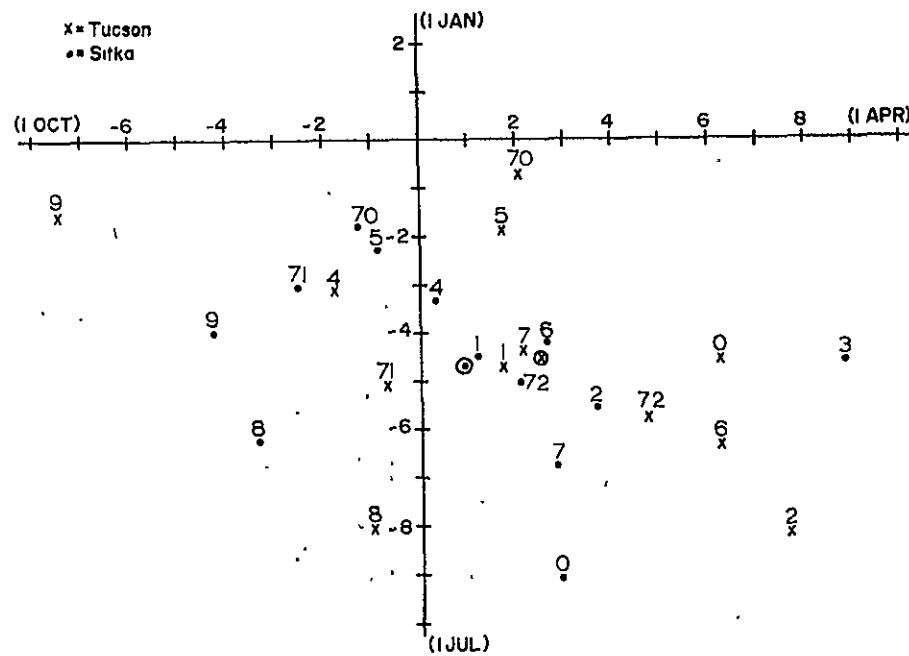
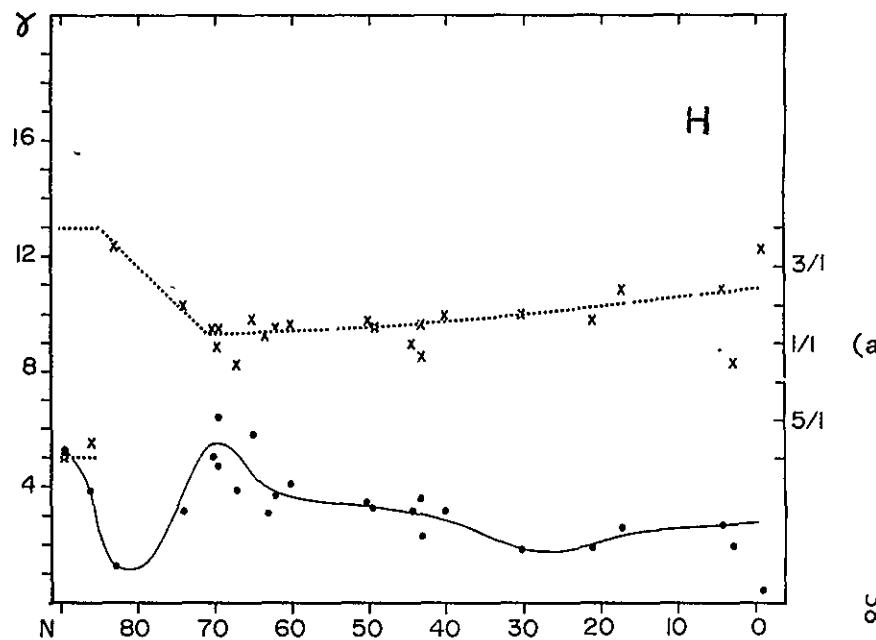
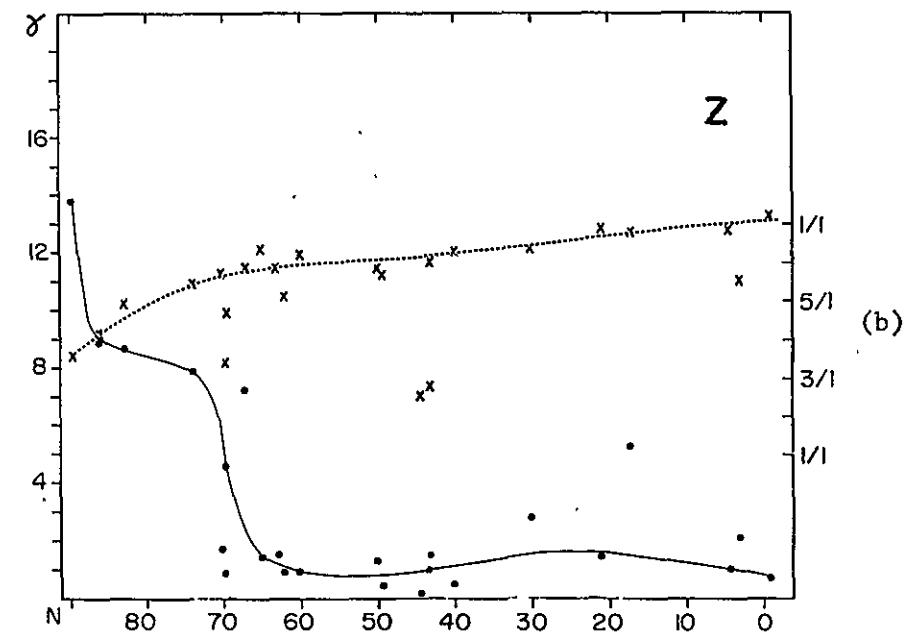


Figure 2. Harmonic dial of the amplitude and phase of the annual wave for each year, 1960-1972, at Sitka and Tucson. The average values for all years are circled. The small number above each point is the year that the point represents (i.e., 3 is for 1963). Axes are labeled in gammas.

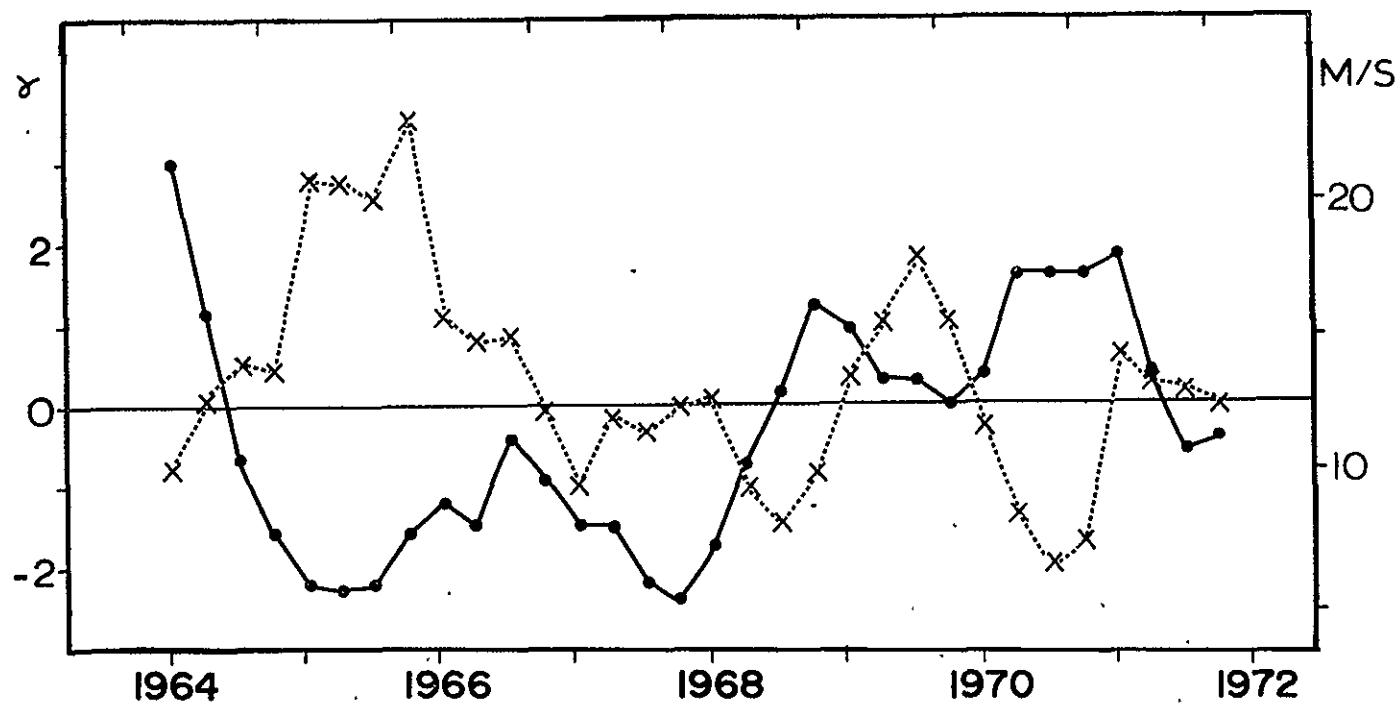


38



(b)

Figure 3. As in Figure 1 except for the semi-annual wave.



65

Figure 4. Twelve-month running mean values of Z (solid line) at Honolulu relative to the parabolic secular trend line, and twelve-month running mean values of zonal wind at 56 km at Barking Sands (dotted line). Tick marks on abscissa are for 1 July; every third month is plotted.

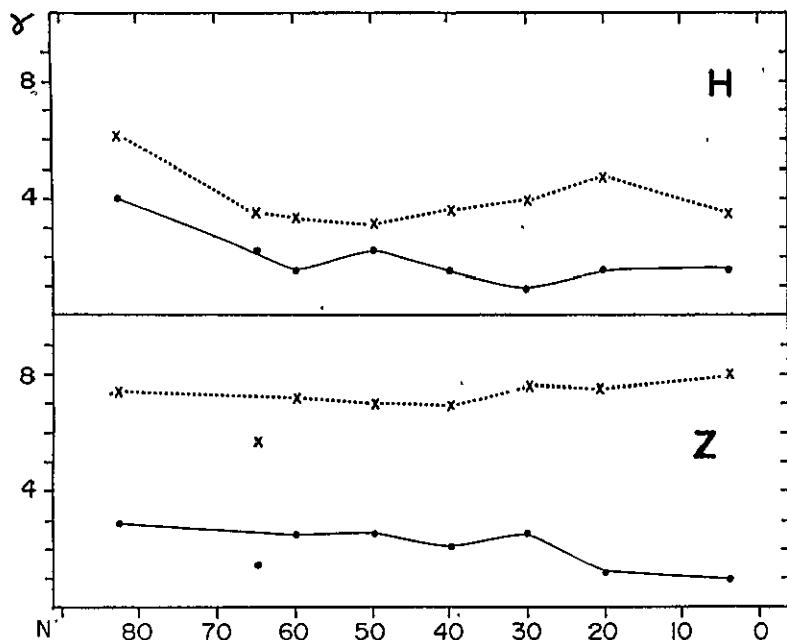


Figure 5. As in Figure 1 except for the QBO. θ scale is months relative to 1 January 1960.

Figures 6a-6f. Height-latitude sections of relative phase lags of MRN and geomagnetic periodic waves in geographic coordinates for station pairs in Table 1. In dotted areas phase lag is within 30° of zero; in shaded regions it is within 30° of 180° . (a) annual wave, U and H (b) semiannual wave, U and H, the dashed line is explained in the text (c) annual wave, U and Z (d) semiannual wave, U and Z (e) annual wave, T and H (f) annual wave, T and Z.

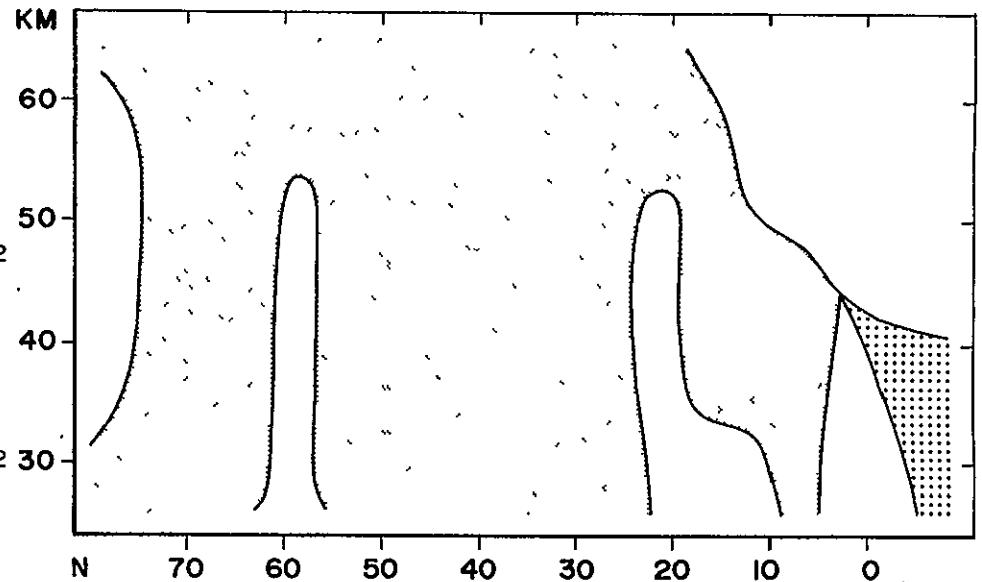


Figure 6a.

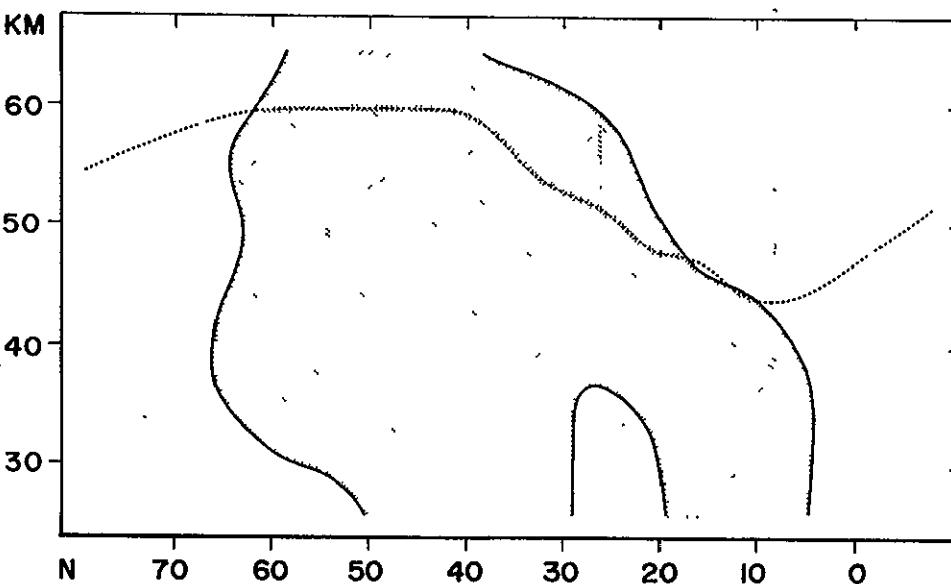


Figure 6b.

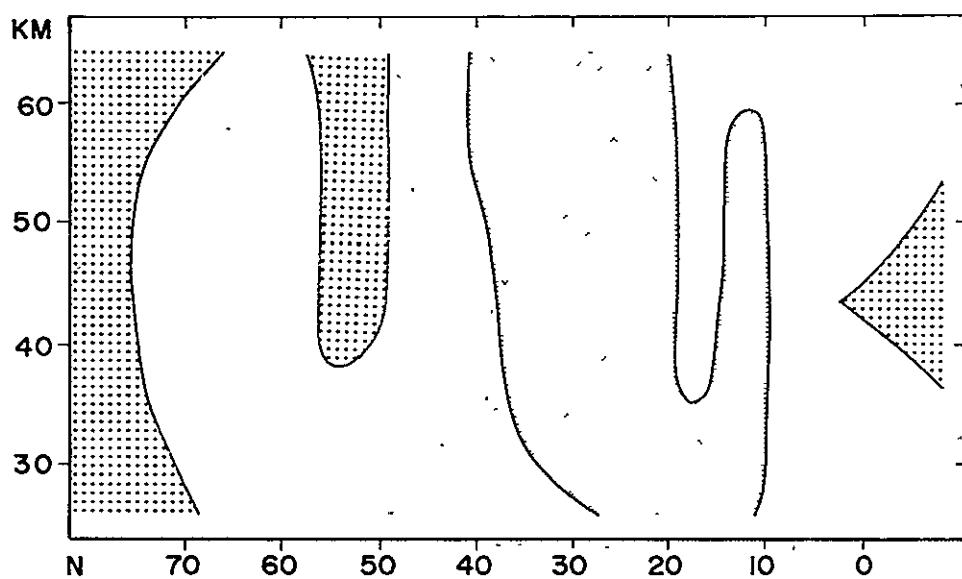


Figure 6c.

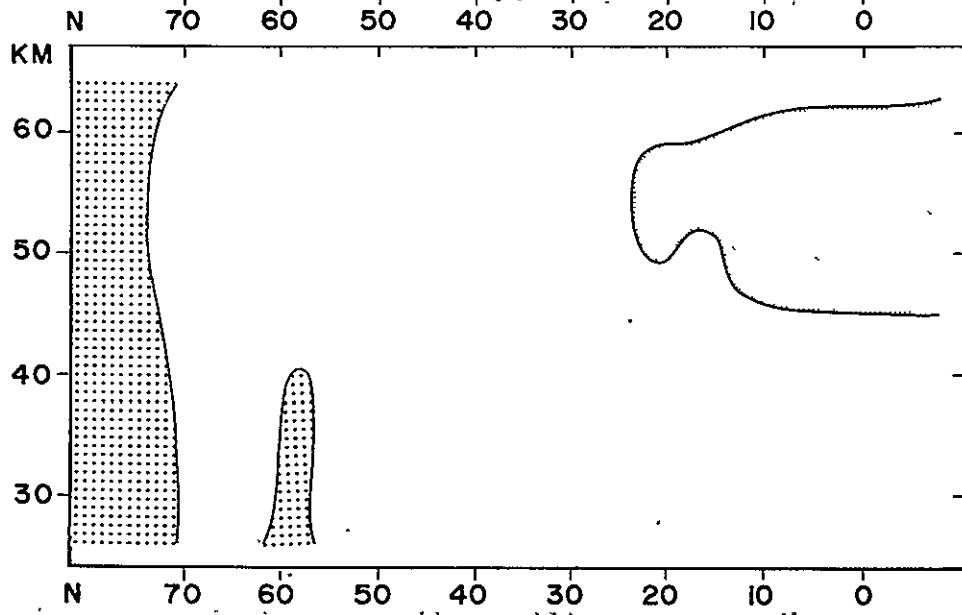


Figure 6d.

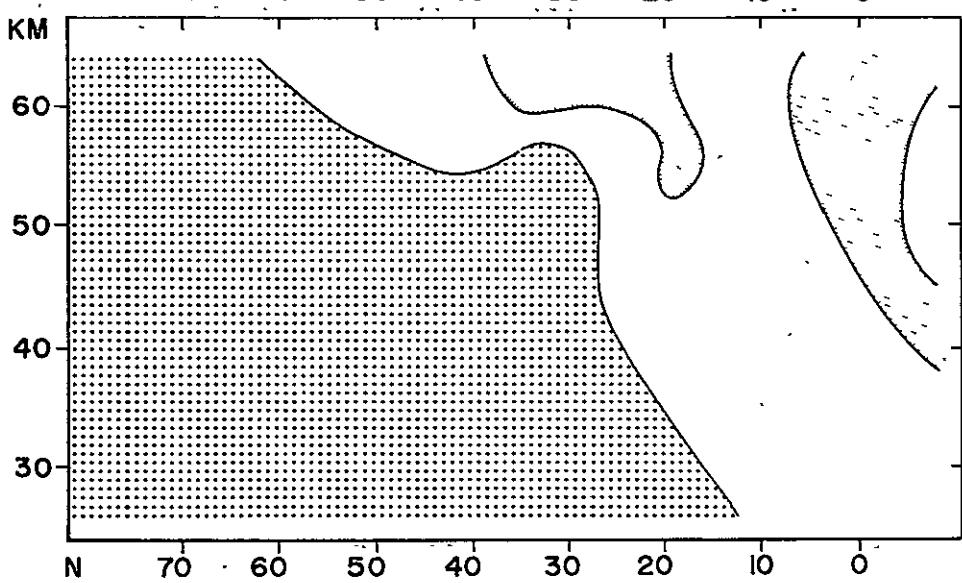


Figure 6e.

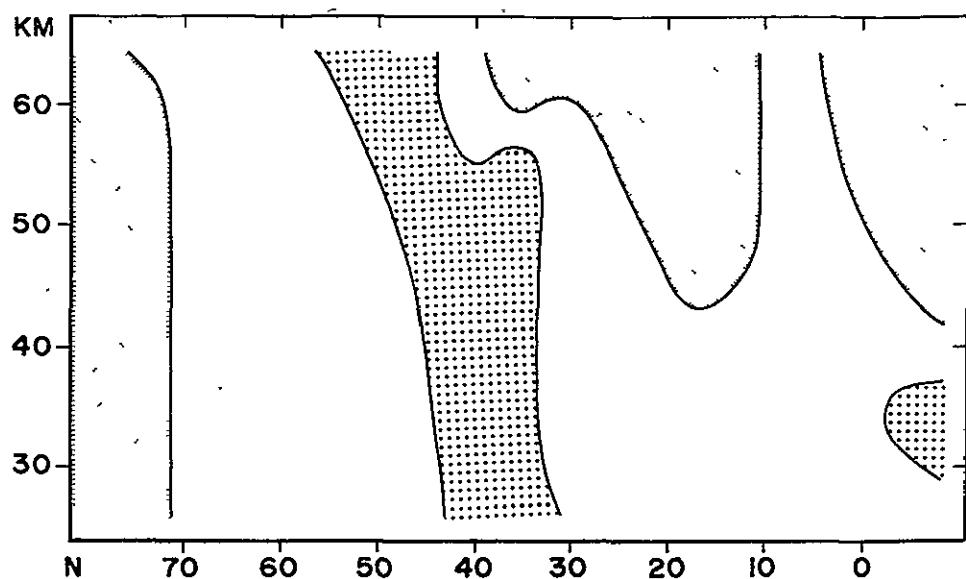


Figure 6f.

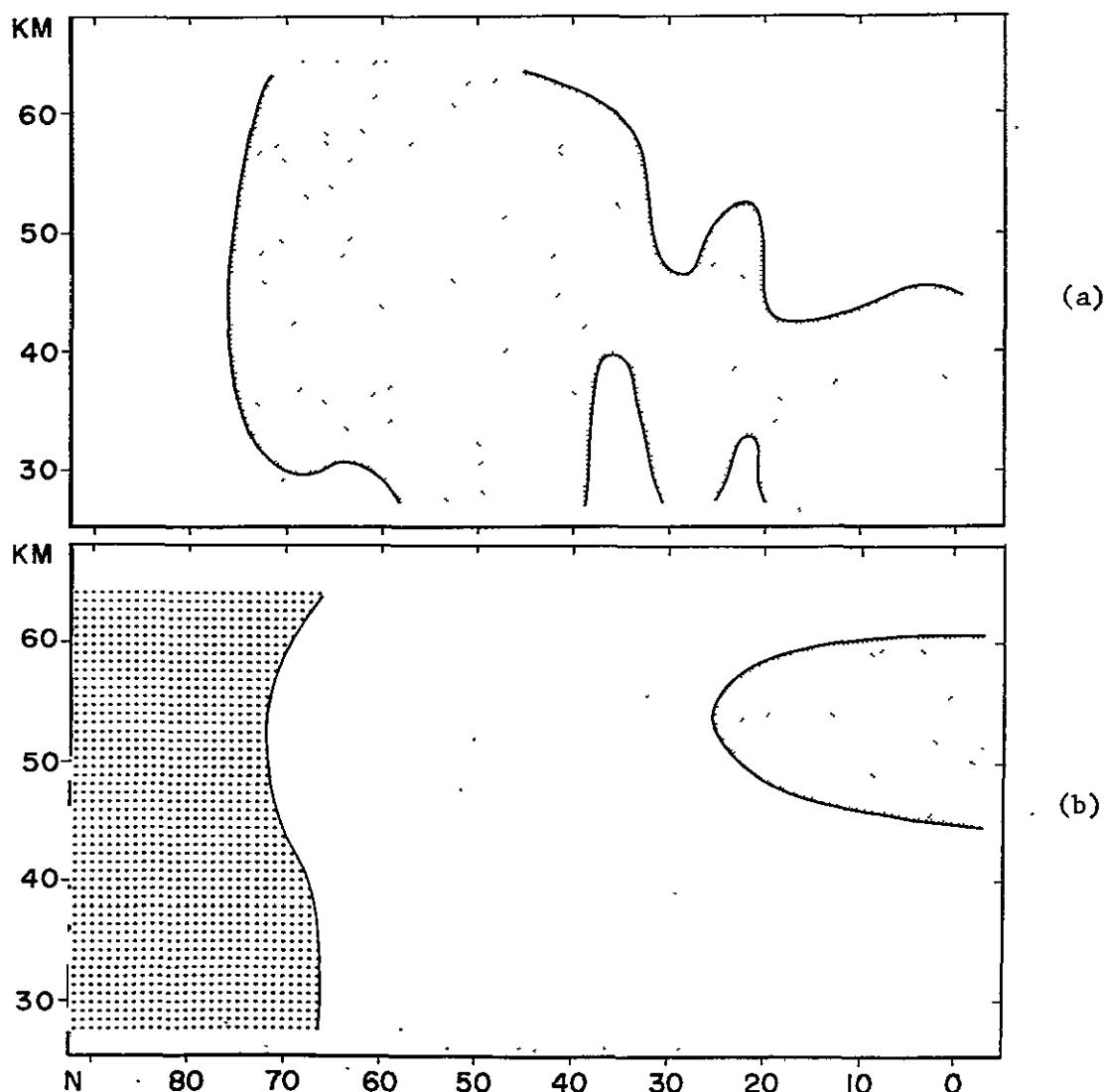
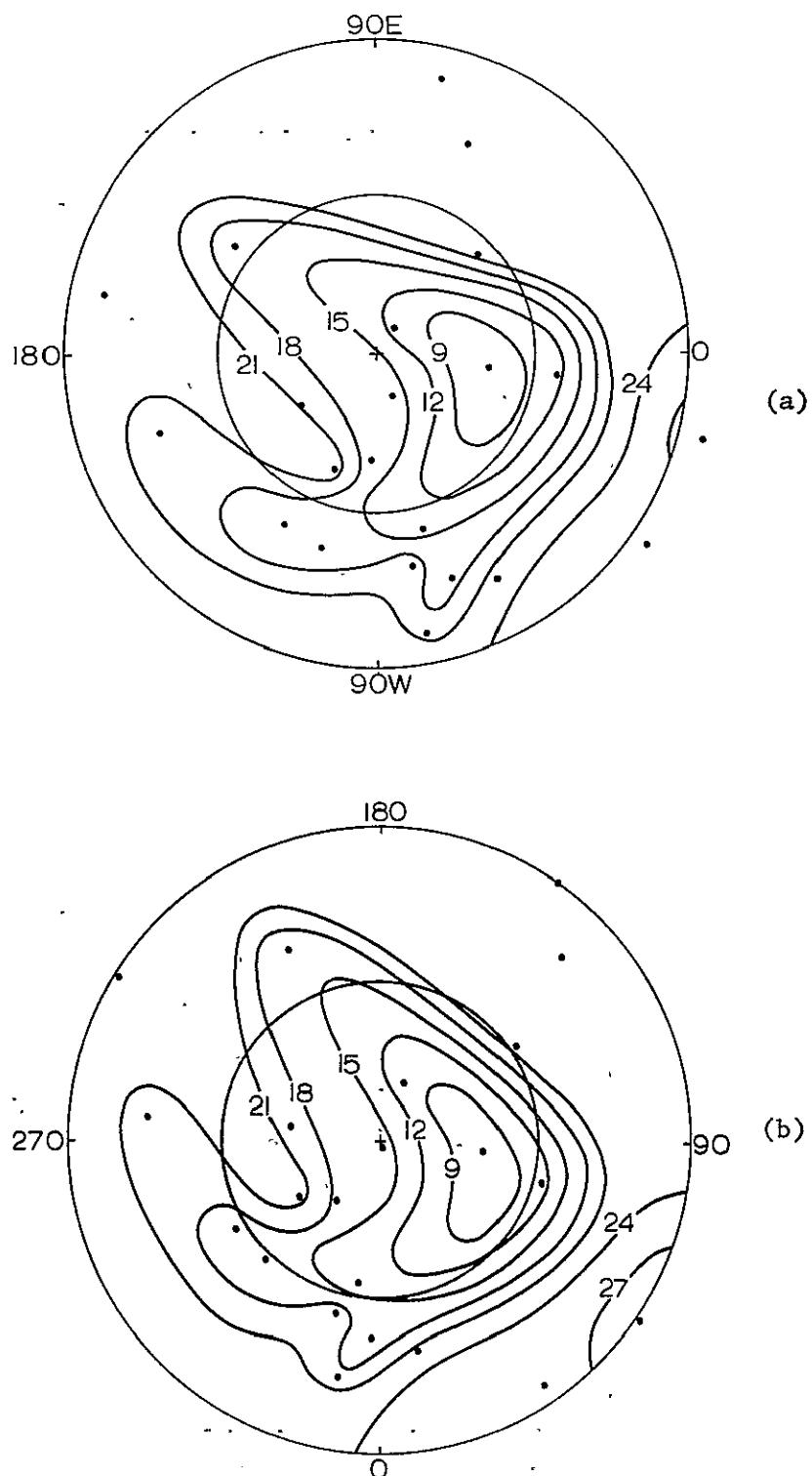


Figure 7. As in Figure 6 except geomagnetic latitude.
 (a) semiannual wave, U and H ; (b) semiannual wave, U and Z .



Figures 8a and 8b. The amplitude (m s^{-1}) of the semiannual wave in zonal wind at the altitude indicated by the dashed line in Figure 6(b) or Table 1. Dots are MRN station locations. (a) geographic coordinates; (b) geomagnetic coordinates.

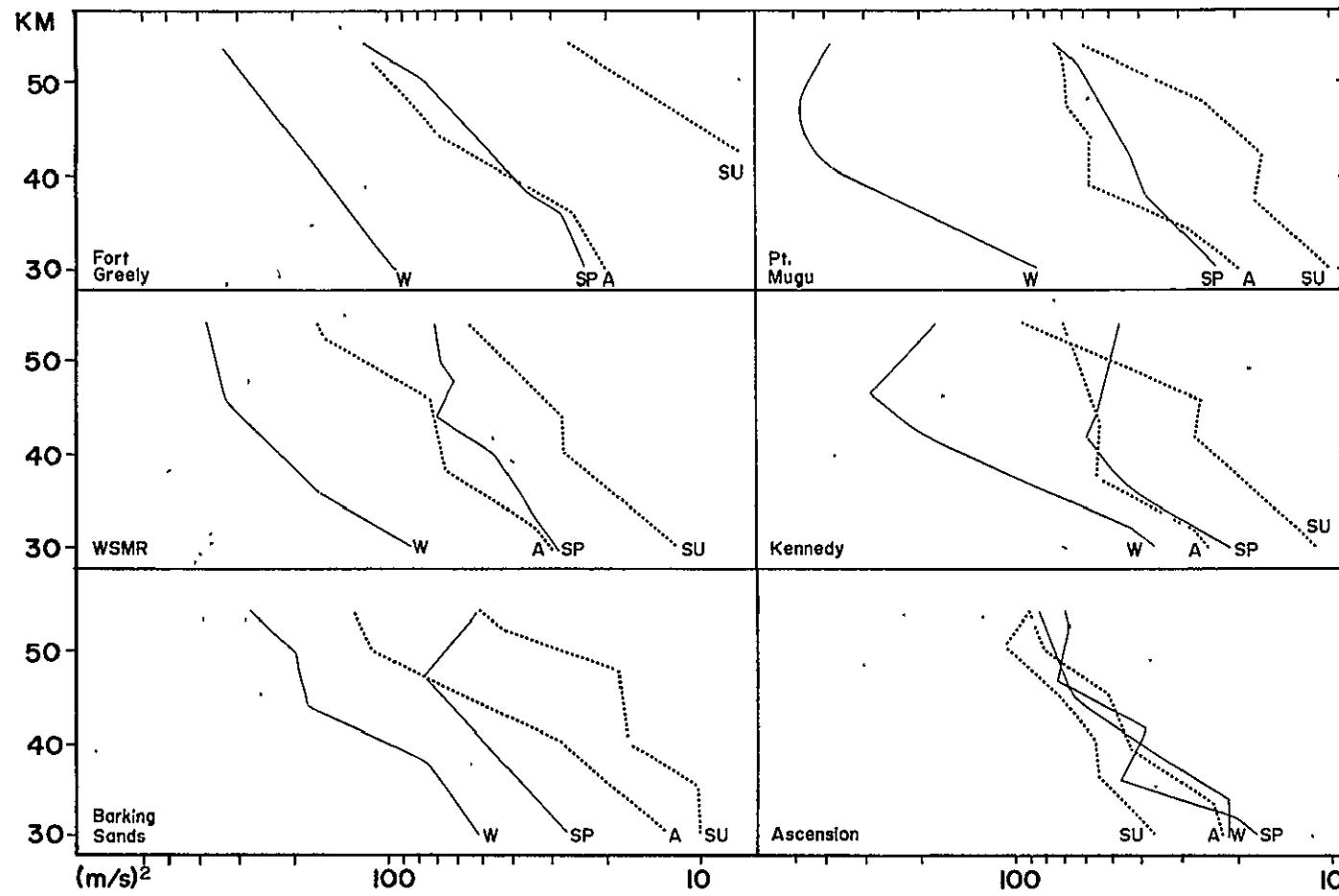


Figure 9. Power spectral density in zonal wind as a function of height for the band centered at $2\pi/11$ days, by season.

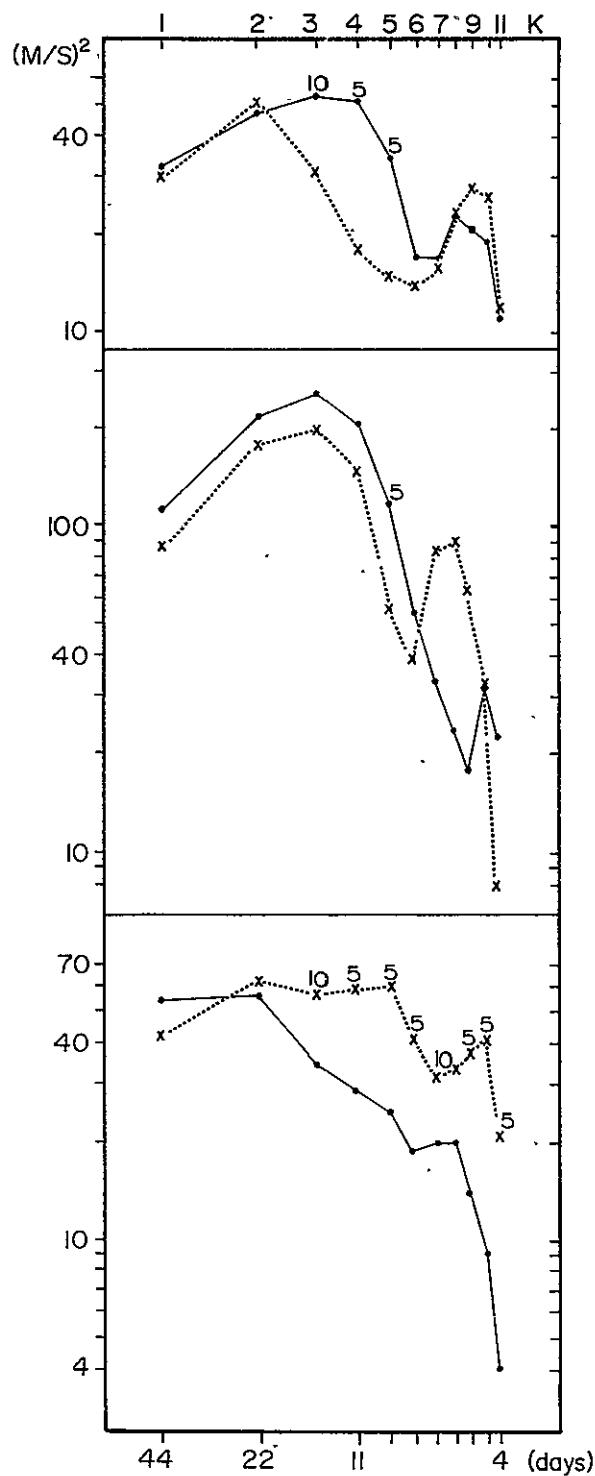


Figure 10. Power spectrum of zonal wind variations at 40 km at Fort Greely. Solid line is for SW years, dotted line is for MSW years, as defined in the text. Statistical significance level of the difference between the two spectra is indicated for continuum values.
 (a) autumn (b) winter (c) spring.

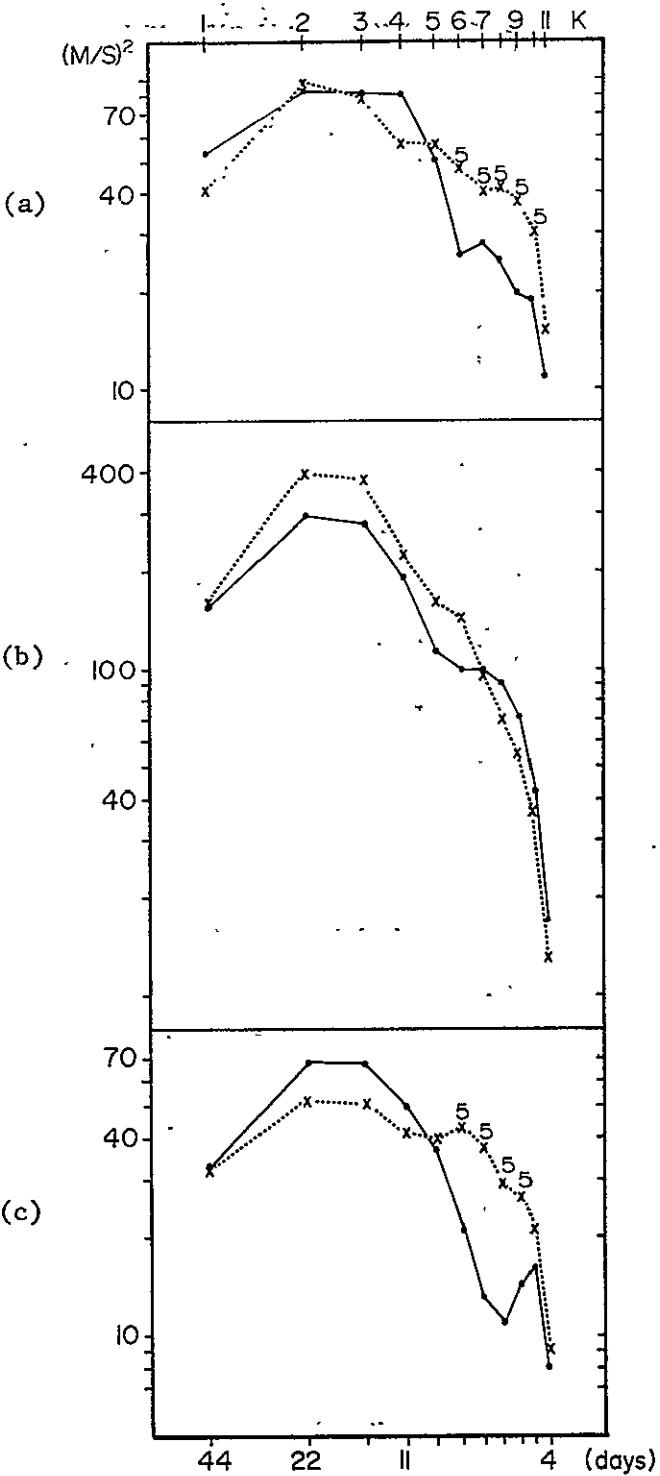


Figure 11. As in Figure 10 except at White Sands.

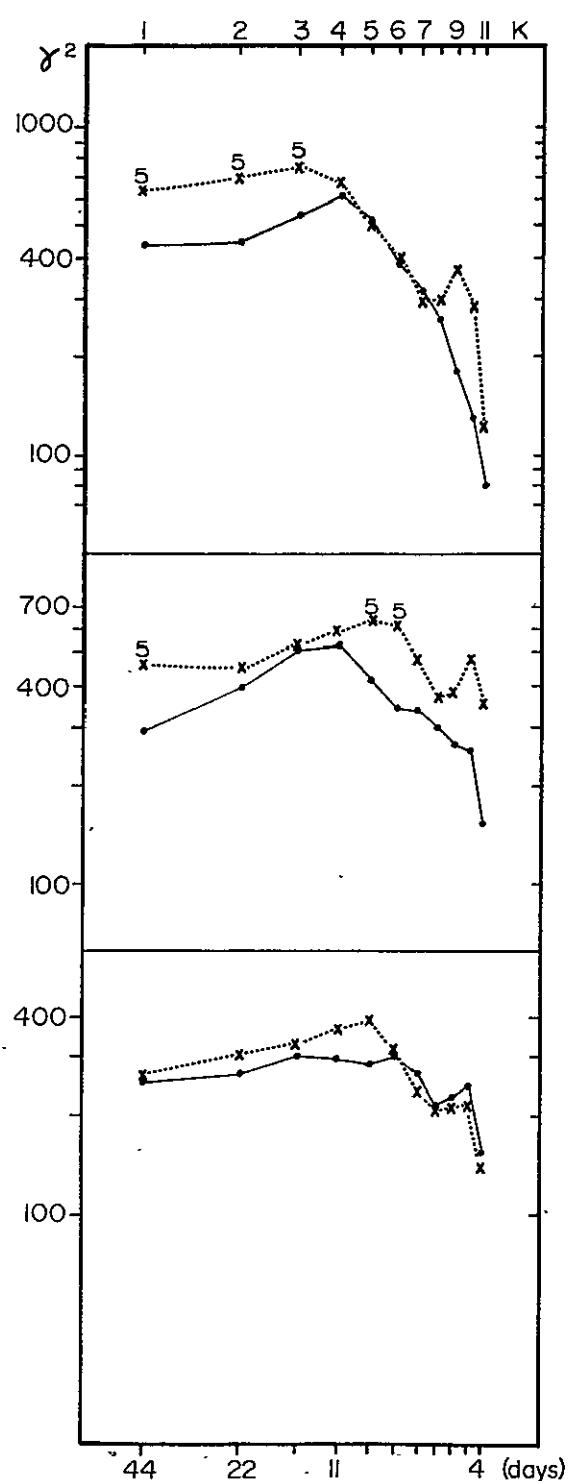


Figure 12. As in Figure 10 except for the horizontal field intensity at College.

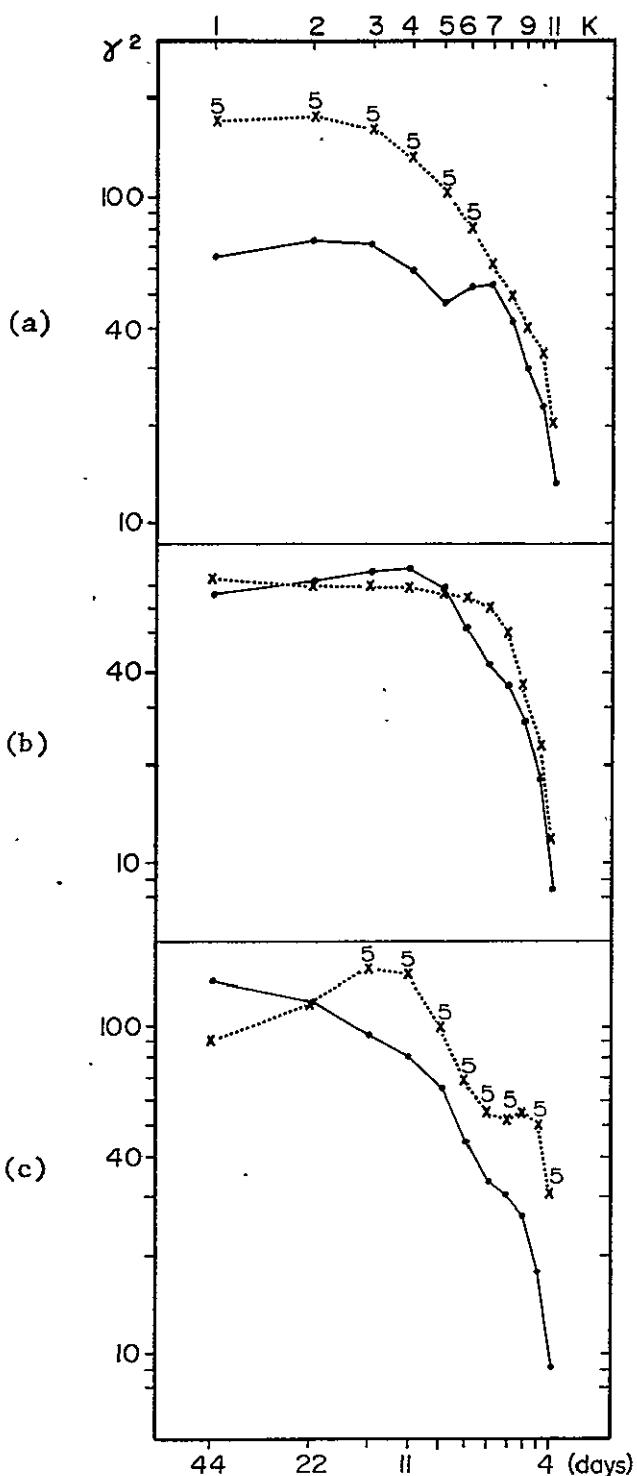


Figure 13. As in Figure 10 except for the horizontal field intensity at Tucson.

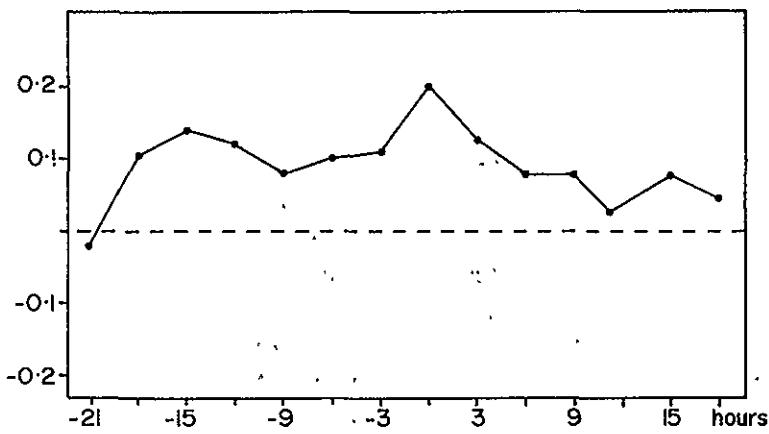


Figure 14. Linear correlation coefficient between layer mean temperature, 40-50 km, at Fort Churchill and K_p as a function of the lag of temperature.

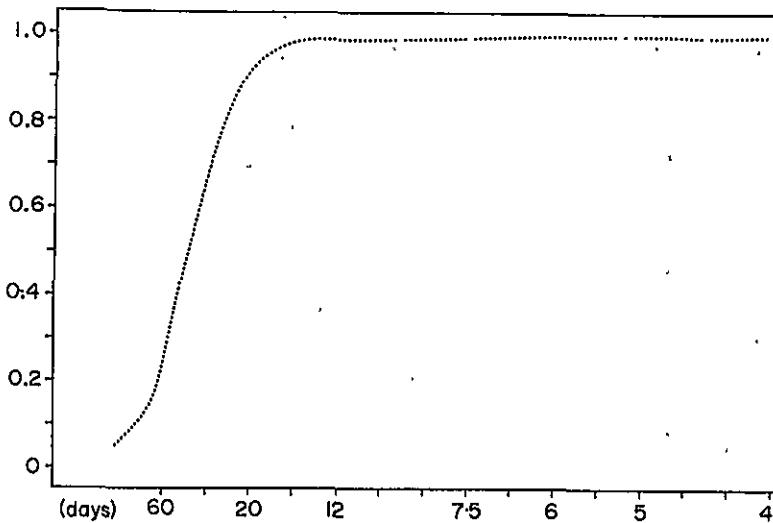


Figure A-1. Theoretical frequency response of the numerical filter described in the Appendix.

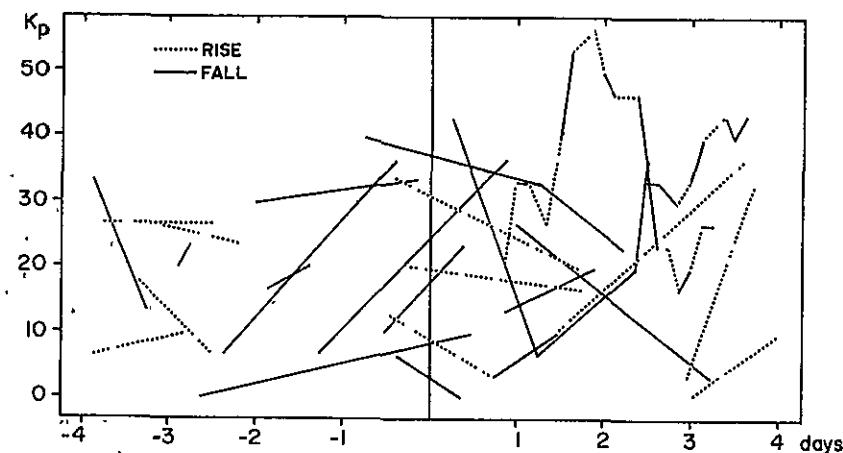


Figure 15. Trend of the temperature at 40 km at Fort Churchill between closely spaced ascents made near a solar sector boundary crossing. End points of each line segment are plotted at the time relative to boundary crossing and at the corresponding value of K_p .

Table 1. List of meteorological rocket stations.

STATION	LAT. (GEOGRAPHIC)	LONG. (GEOGRAPHIC)	LAT. (GEOMAGNETIC)	LONG. (GEOMAGNETIC)	NUMBER OF OBS.	AT 50 KM (WIND)	NEAREST GEOMAGNETIC OBSERVATORY	SYMBOL
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a. Stations used in Figures 6-8.								
THULE	77	69	88	10	335	296	RESOLUTE	RB
FORT GREELY	64	146	64	261	1011	563	COLLEGE	CO
CHURCHILL	59	94	68	324	991	884	CHURCHILL	FC
PRIMROSE LAKE	55	110	62	305	316	312	SITKA	SI
WALLOPS	38	75	48	351	1351	674	FREDRICKSBURG	FR
POINT MUGU	34	119	41	302	1971	1226	BOULDER	BD
WHITE SANDS	32	106	42	317	2481	988	TUCSON	TU
KENNEDY	28	81	38	347	1916	1142	DALLAS	DS
BARKING SANDS	22	160	21	265	1372	898	HONOLULU	HO
ANTIGUA	17	62	28	10	466	371	SAN JUAN	SJ
SHERMAN	9	80	20	350	631	422	FUQUENE	FQ
KWAJALEIN	9	-168	1	238	318	305	GUAM	GU
ASCENSION	-8	14	-1	55	1196	937	HUANCAYO	HU

b. Stations used only in Figure 8.								
HEISS ISLAND	81	-58	72	156	156	(56 KM)		
WEST GEIRNISH	57	7	60	84	124	(56 KM)		
VOLGOGRAD	49	-44	43	125	87	(52 KM)		
RYORI	39	-142	29	207	32	(48 KM)		
ARENOSILLO	37	7	41	76	80	(54 KM)		
SONMIANI	25	-67	16	137	54	(50 KM)		
GRAND TURK	21	71	32	357	170	(50 KM)		
THUMBA	9	-77	0	146	145	(50 KM)		
NATAL	-6	-35	5	34	131	(46 KM)		

b. Stations used only in Figure 8.								
HEISS ISLAND	81	-58	72	156	156	(56 KM)		
WEST GEIRNISH	57	7	60	84	124	(56 KM)		
VOLGOGRAD	49	-44	43	125	87	(52 KM)		
RYORI	39	-142	29	207	32	(48 KM)		
ARENOSILLO	37	7	41	76	80	(54 KM)		
SONMIANI	25	-67	16	137	54	(50 KM)		
GRAND TURK	21	71	32	357	170	(50 KM)		
THUMBA	9	-77	0	146	145	(50 KM)		
NATAL	-6	-35	5	34	131	(46 KM)		

* Minus is south or east.

TABLE 2. Periodic analysis results of the geomagnetic field elements. Amplitudes are in tenths of gammas and phases are in degrees. Statistical errors are in parentheses.

	GEO-MAGNETIC		NUMBER OF MONTHS		H						Z					
					QBO		ANNUAL		SEMIANNUAL		QBO		ANNUAL		SEMIANNUAL	
	LAT	LON	H	Z	AMP	PHASE	AMP	PHASE	AMP	PHASE	AMP	PHASE	AMP	PHASE	AMP	PHASE
THULE	89	358	34	34	46(25)	-78(44)	47(17)	-177(23)	53(17)	-175(20)	190(69)	-102(36)	126(42)	-151(20)	138(40)	151(18)
ALERT	86	168	77	77	46(16)	106(22)	36(15)	21(27)	39(15)	-154(25)	15(24)	84(9)	156(23)	4(8)	89(22)	-173(15)
RESOLUTE	83	168	115	115	40(5)	159(7)	29(5)	-62(10)	13(5)	153(23)	29(13)	-118(30)	192(14)	1(4)	87(13)	-125(9)
BAKER LAKE	74	315	101.	101	89(13)	-15(9)	169(14)	174(5)	32(13)	55(27)	109(20)	134(11)	112(20)	2(11)	79(20)	-94(15)
LIERVOGUR	70	71	87	81	1(13)	-122(95)	12(14)	27(23)	50(4)	20(15)	18(14)	-7(14)	21(14)	23(11)	17(4)	-80(14)
CHURCHILL	69	323	69	69	86(14)	79(8)	102(12)	-151(6)	47(11)	-8(14)	28(12)	81(26)	97(11)	54(7)	9(8)	141(73)
BARROW	69	241	95	94	29(9)	-107(17)	94(8)	-167(5)	64(8)	21(7)	16(6)	150(25)	36(6)	142(10)	46(6)	-141(8)
GREAT WHALE RIV	67	347	64	64	139(16)	30(7)	51(5)	-74(19)	39(15)	-33(25)	40(9)	-43(13)	41(8)	-135(12)	73(9)	-66(7)
COLLEGE	65	257	156	156	22(5)	-43(15)	53(6)	164(6)	58(6)	38(5)	15(3)	112(13)	19(3)	119(10)	14(3)	-40(14)
LERWICK	63	89	20	20	24(10)	81(17)	21(16)	-164(26)	31(5)	10(10)	13(16)	87(20)	17(4)	7(12)	15(3)	-68(12)
MEANOOK	62	301	68	68	7(4)	30(39)	44(4)	-168(5)	37(4)	25(6)	14(4)	-35(15)	15(4)	6(14)	9(3)	-115(25)
SITKA	60	275	156	156	15(4)	-49(15)	48(4)	168(4)	41(4)	30(5)	25(4)	-137(5)	7(2)	21(21)	9(2)	-49(16)
FREDRICKSBURG	50	350	156	156	22(5)	-70(15)	64(5)	163(5)	34(5)	37(9)	26(5)	-148(13)	3(4)	-115(86)	13(5)	-69(25)
BOULDER	49	317	72	72	19(5)	-11(17)	54(5)	-174(6)	33(5)	27(9)	17(3)	-140(12)	12(3)	172(16)	4(3)	-80(61)
STEPANORKA	44	111	32	32	9(8)	143(70)	40(9)	-175(14)	32(9)	-2(16)	47(6)	104(9)	8(5)	78(55)	2(4)	93(87)
CASTLE ROCK	43	299	33	33	~ 19(6)	-88(19)	34(6)	-170(11)	36(6)	-18(10)	9(3)	101(19)	6(3)	-141(33)	15(3)	108(11)
DALLAS	43	328	99	99	15(5)	-94(20)	51(5)	-168(5)	23(5)	29(12)	28(3)	-96(6)	17(3)	167(10)	10(3)	-62(18)
TUCSON	40	312	156	153	15(4)	-39(18)	53(4)	152(5)	32(4)	44(8)	21(3)	-158(8)	13(3)	-114(4)	5(3)	-43(44)
SAN JUAN	30	3	156	156	9(6)	-7(69)	38(4)	175(7)	18(4)	43(14)	26(4)	-108(9)	40(7)	-138(10)	28(7)	-37(15)
HONOLULU	21	266	144	144	15(4)	49(15)	35(4)	143(6)	19(4)	37(12)	12(2)	-116(9)	25(2)	-169(4)	14(2)	-7(8)
FUQUENE	17	355	49	38	8(5)	126(42)	5(4)	167(66)	26(6)	85(12)	81(18)	52(13)	57(16)	-83(37)	53(26)	-14(28)
GUAM	4	213	154	147	15(4)	-45(18)	29(4)	138(9)	27(4)	84(10)	10(3)	-75(16)	46(3)	-174(4)	11(3)	-9(15)
MUNTINLUPA	3	190	60	60	84(7)	50(5)	18(7)	-172(23)	20(7)	-31(21)	19(4)	-48(13)	32(4)	-162(7)	22(4)	-88(11)
HUANCAYO	-1	354	51	44	38(6)	69(10)	24(6)	-138(16)	4(4)	149(77)	4(3)	-13(58)	33(3)	176(6)	7(3)	13(38)
TOOLANGI	-47	221	57	0	3(5)	136(90)	46(7)	24(8)	34(7)	50(11)	-	-	-	-	-	-
ARGENTINE IS.	-54	3	96	96	10(3)	52(22)	36(3)	32(5)	33(3)	50(6)	10(3)	13(18)	26(3)	20(7)	21(3)	65(8)
KERGUELEN	-57	128	55	0	24(6)	-19(16)	26(6)	2(14)	34(6)	17(11)	-	-	-	-	-	-
BYRD	-71	336	91	90	38(37)	62(5)	115(35)	95(19)	15(25)	-100(89)	21(21)	61(5)	73(19)	72(16)	64(19)	2(19)

Table 3. Coherence-squared statistics between MRN and geomagnetic monthly data, and number of data pairs at lag one month. Station symbols as in Table 1.

STATION	30 KM					48 KM					56 KM				
	CO	FC	FR	TU	HO	CO	FC	FR	TU	HO	CO	FC	FR	TU	HO
N(T)	95	65	100	134	87	86	65	89	130	85	77	65	74	114	77
N(U)	114	66	144	148	115	110	66	144	147	117	102	61	131	146	110
a. H-T	.12	.50	.23	.50*	.12	.21	.54	.15	.19	.02	.16	.56	.16	.20	.08
Z-T	.05	.51	.06	.18	.46	.14	.55	.12	.05	.08	.09	.55	.09	.10	.31
H-U	.16	.32	.48*	.52*	.27	.20	.59	.51*	.49*	.27	.15	.53	.48*	.51*	.28
Z-U	.03	.37	.00	.13	.53	.02	.53	.00	.15	.57*	.02	.55	.00	.16	.56*
b. H-T	.39	.28	.05	.11	.13	.19	.18	.12	.05	.18	.11	.10	.16	.11	.14
Z-T	.18	.05	.19	.17	.34	.32	.18	.17	.03	.08	.05	.05	.36	.13	.02
H-U	.05	.42	.42	.40	.09	.20	.29	.39	.50*	.22	.21	.23	.40	.57*	.27
Z-U	.15	.25	.17	.07	.48	.16	.03	.07	.04	.53	.20	.10	.02	.03	.66*
c. H-T	.19	.38	.20	.00	.02	.27	.22	.05	.17	.08	.07	.14	.07	.13	.02
Z-T	.01	.03	.03	.00	.17	.04	.03	.03	.17	.02	.03	.05	.15	.08	.06
H-U	.01	.04	.03	.12	.31	.01	.09	.03	.07	.05	.02	.03	.02	.07	.09
Z-U	.20	.23	.02	.01	.02	.02	.13	.03	.04	.11	.01	.04	.06	.10	.24
d. H-T	.02	.28	.11	.20	.06	.06	.36	.02	.18	.01	.14	.24	.01	.13	.20
Z-T	.22	.36	.03	.07	.07	.13	.45	.05	.10	.08	.01	.38	.08	.10	.14
H-U	.05	.12	.08	.13	.16	.08	.32	.09	.11	.03	.06	.37	.09	.10	.04
Z-U	.05	.38	.02	.03	.07	.02	.35	.00	.03	.27	.01	.41	.01	.03	.15

CODE: a. Annual Wave
 b. Semiannual Wave
 c. Terannual Wave
 d. Quasi-biennial Wave

STATISTICAL SIGNIFICANCE INDICATORS: * Value exceeds 0.1% confidence level.
 = Value exceeds 1% confidence level.
 - Value exceeds 5% confidence level.

Station codes are given in Table 1.

"H-T" = Comparison of horizontal component of the geomagnetic field strength with temperature at specified level from nearest rocket observation.

Table 4a. Linear correlation coefficients of monthly values of MRN and geo-magnetic data. Those which meet the 1% significance level are asterisked, 5% level are underlined.

PARAMETERS	(LEVEL)	CO	FC	FR	TU	HO
T-H	30	.083	.491*	<u>.223</u>	.280*	-.138
	48	.121	.485*	.014	.247*	-.070
	56	.145	.363*	-.127	.161	.071
U-H	30	-.188	-.354*	-.313*	-.218*	-.211
	48	-.277*	-.488*	-.350*	-.320*	-.265*
	56	-.338*	-.466*	-.343*	-.351*	-.222
T-Z	30	.230	-.280	.117	-.026	.293*
	48	.253	<u>-.301</u>	.089	<u>-.181</u>	-.143
	56	.118	-.309	-.164	<u>-.246</u>	-.303*
U-Z	30	-.020	<u>.286</u>	.064	-.110	-.282*
	48	-.026	<u>.284</u>	.047	-.103	-.490*
	56	.010	<u>.303</u>	.051	-.094	-.504*

Table 4b. Same as above except the annual waves were first subtracted from the data. Note the lack of significant correlation here compared with above.

T-H	30	-.114	.174	.111	.099	-.227
	48	-.195	.136	.034	.165	-.044
	56	-.024	.006	-.090	<u>.195</u>	.173
U-H	30	.013	-.033	-.077	-.038	-.077
	48	-.149	-.226	-.083	-.105	-.051
	56	-.190	-.200	-.020	-.109	.053
T-Z	30	.169	-.002	.132	-.024	.008
	48	.154	-.048	.111	-.167	-.102
	56	-.027	-.151	-.015	-.184	-.172
U-Z	30	.095	.101	.116	-.026	-.189
	48	.083	.086	.020	-.063	-.203
	56	.113	.157	-.014	-.082	-.196

TABLE 5. Yearly values of solar and geophysical parameters. Some relative maxima are underlined; for annual wave relative minima are underlined.

YEARS	61	62	63	64	65	66	67	68	69	70	71
SUN SPOT NO.	54	38	28	10	15	47	94	106	<u>106</u>	105	67
10.7 cm FLUX	104	84	80	72	76	103	143	149	151	<u>156</u>	113
<u>SEMIANNUAL AMPLITUDES</u>											
SITKA (H)	4.7	4.3	4.0	3.2	2.3	4.1	2.7	3.7	<u>6.9</u>	5.5	4.1
FREDRICKSBURG (H)	2.9	3.9	5.0	4.1	1.6	4.5	2.8	1.9	<u>6.9</u>	5.5	3.4
TUCSON (H)	3.3	3.1	4.4	4.0	2.3	4.5	4.8	2.6	<u>6.1</u>	4.5	2.6
GREELY (U - 48KM)					9.2	9.2	12.8	19.5	<u>21.5</u>	11.9	
WALLOPS (U - 48KM)	23.7		24.0			9.9	16.8	8.5	8.7	<u>18.4</u>	16.4
MUGU (U - 48KM)		16.4	14.8			10.0	4.4	11.5	13.5	<u>20.0</u>	17.4
WSMR (U - 48KM)	11.0	19.5	15.5	15.1	8.0	8.8	11.6	13.4	<u>24.6</u>	18.8	11.3
BARKING SANDS (U - 48KM)					19.3				23.9	<u>26.5</u>	25.0
<u>ANNUAL AMPLITUDES</u>											
GREELY (U - 48KM)					21.3	25.5	32.9	27.1	<u>13.2</u>	28.6	
WALLOPS (U - 48KM)	49.8		43.0		74.5	60.0	57.6	61.5	<u>59.1</u>	65.8	61.9
MUGU (U - 48KM)		53.9	46.0		63.6	60.4	49.1	51.8	<u>46.8</u>	55.6	53.4
WHITE SANDS (U - 48KM)	54.4	50.3	46.0	58.3	63.8	56.7	48.0	49.0	<u>43.5</u>	48.4	52.4
BARKING SANDS (U - 48KM)					42.4				36.1	<u>37.6</u>	27.9
											38.2

(a) SEASON= WINTER POWER (M2/SEC2)										SEASON= SPRING POWER (M2/SEC2)										SEASON= SUMMER POWER (M2/SEC2)										SEASON= AUTUMN POWER (M2/SEC2)									
VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.										
KH *										*																													
60 *	80	303	534	406	236	283	377	397	305	231	118	497	3	159330	*	11	26	51	58	55	42	48	71	74	90	58	61	11	135135										
60 *	60	301	454	335	243	324	401	375	272	203	104	443	11	135135	*	4	41	66	69	68	55	53	74	83	96	59	94	34	96	46									
59 *	59	341	411	281	237	314	362	335	234	157	81	367	30	102102	*	11	70	99	97	88	69	66	85	92	92	50	117	73	51	51									
59 *	59	341	411	281	237	314	362	335	234	157	81	367	30	102102	*	24	84	118	116	92	69	73	87	84	74	38	132	111	21	21									
110 *	110	436	500	516	203	287	244	264	206	139	71	397	61	64	*	34	81	192	100	82	66	66	71	67	60	30	105	139	21	21									
165 *	165	485	516	319	185	149	165	193	159	119	61	356	102	24	*	38	70	75	76	75	67	59	59	57	46	21	91	153	20	20									
184 *	184	420	402	275	182	124	137	155	118	84	40	316	115	23	*	35	55	58	67	72	65	57	59	54	42	20	82	164	20	20									
166 *	166	326	287	225	164	92	109	126	67	52	22	223	123	23	*	35	48	50	62	68	60	55	59	52	44	23	81	172	20	20									
130 *	130	272	285	227	139	71	74	81	62	48	24	219	133	22	*	41	54	49	55	58	50	48	51	43	39	21	76	182	19	19									
116 *	116	242	273	214	149	49	62	69	55	46	24	189	138	22	*	49	61	49	48	47	36	36	37	31	30	16	62	188	19	19									
60 *	90	197	226	176	86	46	61	59	42	31	15	140	144	22	*	50	59	44	42	40	30	26	28	23	11	57	189	19	19										
92 *	92	182	195	140	78	66	52	37	26	9	137	145	22	*	48	53	37	35	33	24	22	26	20	9	48	192	19	19											
88 *	88	176	177	119	77	70	70	49	34	33	12	129	145	21	*	46	51	38	25	25	17	18	25	26	8	45	192	19	19										
85 *	85	166	164	116	83	72	67	45	34	33	14	132	143	22	*	45	52	34	26	22	16	17	21	23	8	41	194	19	19										
88 *	88	166	153	106	79	65	52	35	27	27	12	116	145	21	*	43	51	33	25	21	16	15	16	18	7	43	201	19	19										
30 *	30	92	169	147	95	73	54	43	30	25	23	10	105	146	22	*	42	50	32	23	20	16	13	13	16	15	7	30	204	19	19								
PER.	PER.	44	22	14.7	11	8.8	7.3	6.3	5.5	4.9	4.4	4	(DAYS)	PERIOD	PERIOD	44	22	14.7	11	8.8	7.3	6.3	5.5	4.9	4.4	4	(DAYS)												

(a) SEASON= WINTER POWER (M2/SEC2)										SEASON= SPRING POWER (M2/SEC2)										SEASON= SUMMER POWER (M2/SEC2)										SEASON= AUTUMN POWER (M2/SEC2)									
VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.										
KH *										*																													
60 *	6	30	32	31	33	37	50	65	56	54	25	60	23	105105	*	10	105	194	177	119	91	55	14	34	51	21	137	1	162162										
6	6	24	34	35	34	40	44	53	54	43	19	60	75	49	49	*	8	98	157	135	98	83	63	32	52	80	42	97	13	122122									
5	5	22	32	35	33	35	32	34	32	28	12	41	115	23	23	*	1	81	106	82	79	83	81	70	75	101	61	126	47	82	82								
3	3	17	26	26	25	25	24	23	21	22	12	27	136	22	22	*	4	67	94	85	103	109	98	94	81	103	71	105	77	47	4726								
50 *	50	15	23	21	14	21	21	14	20	22	12	29	142	22	22	*	6	65	107	111	120	117	105	94	69	98	75	174	97	28130									
50 *	1	12	20	17	14	16	17	17	18	16	9	21	145	22	22	*	17	61	100	108	94	86	93	77	47	58	41	91	113	21	86								
50 *	2	0	14	12	11	11	12	12	13	13	6	16	147	22	22	*	28	58	76	80	62	53	63	36	39	35	18	88	120	23	23								
1	1	7	10	10	10	11	11	19	10	19	5	11	151	22	22	*	27	56	68	69	50	37	40	41	35	34	19	58	123	23	23								
1	1	5	9	10	11	11	11	10	9	8	4	14	152	22	22	*	50	64	69	49	33	35	35	27	26	15	68	129	22	22									
1	1	4	7	8	8	8	9	4	7	6	4	9	155	22	22	*	24	47	52	52	48	27	29	30	25	24	14	48	133	22	22								
40 *	1	1	4	5	5	5	5	5	4	5	3	3	6	168	21	99	*	31	49	43	37	25	15	16	23	24	21	11	43	135	22	22							
1	1	4	4	4	4	4	4	3	3	4	2	4	158	22	22	*	33	48	39	32	21	9	9	16	18	16	8	36	137	22	22								
1	1	3	5	5	5	4	4	3	3	3	2	6	156	22	22	*	29	51	32	26	18	9	10	15	15	11	5	29	140	22	22								
1	1	3	4	4	4	3	3	3	3	3	2	3	154	22	22	*	27	36	21	23	17	11	12	15	13	9	3	29	144	21	21								
30 *	30	1	2	3	3	3	2	2	2	1	1	1	2	154	22	22	*	24	30	21	20	16	11	10	11	7	3	21	145	21	21								
PER.	PER.	44	22	14.7	11	8.8	7.3	6.3	5.5	4.9	4.4	4	(DAYS)	PERIOD	PERIOD	44	22	14.7	11	8.8	7.3	6.3	5.5	4.9	4.4	4	(DAYS)												

(b) SEASON= WINTER POWER (M2/SEC2)										SEASON= SPRING POWER (M2/SEC2)										SEASON= SUMMER POWER (M2/SEC2)										SEASON= AUTUMN POWER (M2/SEC2)									
VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.										
KH *										*																													
60 *	4	22	38	36	32	34	46	57	49	39	20	45	23	103191	*	11	60	53	46	88	136	100	28	49	118	77	72	1	156344										
3	3	18	30	28	27	42	43	51	45	38	19	57	75	49	49	*	4	70	79</																				

SEASON= SUMMER												SEASON= AUTUMN																				
POWER (M2/SEC2)												POWER (M2/SEC2)																				
VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.																					
KM	0																															
60	0	14	16	4	27	58	33	12	35	47	30	10	45	8	134212	0	83	128	108	154	136	69	83	142	187	163	68	162	40	95	95	
0	9	17	13	26	42	28	18	35	42	30	12	36	15	127127	0	68	119	112	160	143	75	90	157	202	189	89	215	57	74	74		
0	1	26	26	27	28	23	23	29	30	28	15	35	22	110110	0	52	127	139	164	138	73	87	147	179	187	100	206	68	60	60		
0	4	19	26	23	24	24	24	26	27	16	32	37	91	91	0	57	156	166	162	131	72	64	90	108	125	71	161	80	46	46		
0	5	12	21	18	18	22	24	22	20	22	12	25	50	77	77	0	66	171	166	152	131	87	54	42	59	88	56	150	91	34	34	
50	0	2	10	20	21	19	20	19	16	14	12	6	20	60	56.66	0	75	183	159	120	102	86	52	23	46	89	58	141	99	25114		
0	0	10	19	23	19	16	14	12	10	9	5	20	67	59	59	0	78	182	150	87	57	61	42	18	48	84	51	125	103	24124		
0	1	10	17	18	14	11	11	10	11	6	17	71	54	54	0	68	155	134	77	39	39	36	27	42	51	27	91	105	24	24		
0	1	4	13	12	9	10	11	9	7	8	6	12	74	51	51	0	57	132	117	75	46	36	32	34	36	22	7	83	106	24	24	
0	0	5	9	9	8	9	11	10	7	6	4	11	76	45103	0	53	115	96	62	48	36	22	22	29	22	9	77	108	24	24		
40	0	1	3	6	7	6	6	8	8	7	5	3	8	77	48	48	0	50	97	73	46	39	30	15	11	22	25	12	53	109	24	24
0	2	5	5	4	4	5	5	5	4	4	2	5	77	48	48	0	48	86	58	38	39	29	14	11	20	22	9	55	109	24	24	
0	2	6	5	4	4	5	6	6	6	4	3	2	6	77	47139	0	42	78	52	35	44	36	15	13	21	18	6	50	110	24	24	
0	2	6	5	4	3	5	6	6	5	4	2	7	77	48	48	0	39	76	52	31	42	36	12	13	25	21	8	50	111	24	24	
0	1	5	5	4	4	5	5	5	5	4	2	5	77	48	48	0	36	74	51	25	35	31	9	11	28	28	11	51	111	24	24	
30	0	1	4	4	4	4	5	5	4	4	2	0	77	48	48	0	31	70	48	22	30	28	8	10	28	29	13	40	112	23	23	
PERIOD	44	22	14.7	11	8.8	7.3	6.3	5.5	4.9	4.4	4	(DAYS)				PERIOD	44	22	14.7	11	4.9	7.3	6.3	5.5	4.9	4.4	4	(DAYS)				

SEASON= WINTER POWER (M2/SEC2)										SEASON= SPRING POWER (M2/SEC2)															
		VAR	N	P.R.E.			VAR	N	P.R.E.																
KM *							*																		
60	*	2 223 359 340 305 282 221 172 220 184 50 280 29 111111	*	7 65 70 36 36 41 60 78 56 33 15 76 19 142453																					
	*	4 236 373 395 375 319 256 196 194 156 47 376 44 94 94	*	10 58 60 36 33 40 56 63 46 33 17 62 34 103103																					
	*	14 220 347 430 435 358 292 210 161 126 43 389 51 81 81	*	10 45 49 39 38 45 50 749 43 41 23 52 44 87 87																					
	*	26 204 321 381 370 351 317 211 150 117 40 315 66 62 62	*	6 40 53 49 48 46 40 48 54 53 29 76 61 63141																					
	*	38 219 327 326 281 306 336 251 149 93 32 375 89 36 36	*	6 42 59 56 48 36 34 46 52 53 29 58* 74 51 51																					
50	*	47 204 281 274 231 226 275 256 145 69 27 261 104 24 24	*	10 48 60 57 46 34 36 46 47 42 20 65 90 34 34																					
	*	69 201 239 231 205 171 198 211 122 55 30 250 108 23 23	*	20 62 69 59 43 29 31 41 46 31 9 59 97 28 28																					
	*	87 216 234 201 177 154 161 163 101 56 32 220 113 23 23	*	23 70 84 70 43 24 25 34 39 23 3 62 98 26 26																					
	*	91 219 235 183 159 151 130 111 90 70 34 211 114 23 23	*	16 72 102 86 45 22 24 34 36 18 2 62 99 25 25																					
	*	103 237 243 175 147 145 105 73 78 69 29 199 115 23 23	*	10 75 116 94 42 15 20 33 34 18 4 69 99 25 25																					
40	*	116 254 243 158 125 120 90 77 81 58 21 193 117 23 23	*	6 72 119 95 39 11 14 27 28 17 6 56 98 26 26																					
	*	123 260 248 152 106 99 83 89 89 55 18 185 118 23 23	*	4 72 121 93 34 8 12 29 33 18 4 58 98 26 26																					
	*	122 269 274 167 97 90 74 73 78 48 14 197 119 23 23	*	1 73 126 91 31 7 12 37 45 20 1 62 99 25 25																					
	*	114 270 275 160 90 87 70 56 61 41 10 170 122 23 23	*	0 67 119 85 26 7 11 36 45 18 0 60 99 23105																					
	*	105 265 255 133 84 83 59 46 60 47 13 168 122 23 23	*	1 54 100 74 27 9 8 26 36 19 3 47 99 25 25																					
30	*	101 265 248 122 81 77 48 40 65 54 16 149 122 23 23	*	1 45 87 69 28 12 8 21 31 21 6 43* 99 25 25																					

PER.	44	22	14.7	11	8.8	7.3	6.3	5.5	4.9	4.4	4	(DAYS)	PERIOD	44	22	14.7	11	8.8	7.3	6.3	5.5	4.9	4.4	4	(DAYS)

Table A-2. Analysis of the high frequency variability of the wind at Churchill: (a) Zonal, (b) Meridional.

SEASON= WINTER POWER (M ² /SEC ²)												SEASON= SPRING POWER (M ² /SEC ²)												SEASON= AUTUMN POWER (M ² /SEC ²)														
VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.																		
60	5	79	88	57	105	200	187	73	8	27	7	115	5	230	15	*	11	19	44	52	36	33	51	57	48	42	22	50	3	137256								
60	15	89	112	80	92	153	141	60	16	13	11	89	14	124272	*	11	17	41	49	39	34	44	51	46	42	23	57	21	113113									
60	32	117	158	104	73	114	109	56	48	68	39	145	28	104104	*	8	16	38	46	41	33	32	38	37	34	19	41	39	88	88								
60	29	139	172	108	87	131	118	60	47	71	44	142	42	87	*	4	19	36	40	38	33	29	30	27	24	13	39	65	59	59								
50	*	22	132	151	126	137	150	114	67	56	59	30	135	50	77	77	*	0	24	36	31	31	33	28	26	23	20	11	38	87	35	35						
50	*	9	120	140	136	163	164	122	90	83	66	28	167	58	68	68	*	2	24	37	31	25	25	23	21	18	17	10	30	96	26	26						
50	*	9	118	135	126	129	143	123	92	77	67	34	147	68	57	57	*	2	21	36	35	24	17	17	18	18	18	10	29	103	21	21						
50	*	9	17	107	119	95	91	122	113	73	48	47	27	112	73	52	52	*	4	21	30	30	22	14	14	18	21	23	12	30	107	22	22					
50	*	17	88	95	81	97	121	101	55	27	26	17	105	75	50	50	*	4	19	24	26	25	17	12	18	23	25	14	25	108	21	21						
50	*	13	66	68	71	94	98	72	45	23	19	13	79	73	52	52	*	4	18	26	32	32	21	11	15	20	23	14	34	108	21	21						
50	*	11	50	52	63	81	68	42	33	23	14	9	58	72	53	53	*	3	16	26	32	29	18	11	13	17	20	12	28	107	21	21						
40	*	8	44	49	52	62	54	35	23	20	14	8	50	72	53	53	*	1	12	20	24	22	15	10	13	15	15	8	20	106	21	21						
40	*	6	38	46	45	45	41	31	25	20	16	9	45	72	53	53	*	0	10	16	20	15	10	10	12	11	9	5	18	105	22	22						
40	*	5	32	40	40	38	31	22	21	21	17	8	37	68	57	57	*	0	10	15	18	18	13	8	8	9	8	4	14	105	22	22						
40	*	6	29	34	28	27	24	14	19	18	9	31	68	57	57	*	0	9	13	15	15	12	8	8	8	5	2	13	105	22	22							
30	*	7	28	30	20	19	21	12	11	16	19	11	22	67	59	59	*	1	7	12	15	15	12	7	8	7	3	0	10	103	22	22						
30	*	1	4	7	7	9	5	6	4	5	3	6	6	119	20	20	*	7	31	17	6	29	45	41	31	21	30	23	41	1	156398							
PER.	44	22	14.7	11	8.8	7.3	6.3	5.5	4.9	4.4	4	(DAYS)	PERIOD	44	22	14.7	11	8.8	7.3	6.3	5.5	4.9	4.4	4	(DAYS)	PER.	44	22	14.7	11	8.8	7.3	6.3	5.5	4.9	4.4	4	(DAYS)
SEASON= SUMMER POWER (M ² /SEC ²)												SEASON= AUTUMN POWER (M ² /SEC ²)												SEASON= WINTER POWER (M ² /SEC ²)														
VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.						
60	11	25	59	53	38	45	35	30	73	74	24	67	4	142314	*	7	31	17	6	29	45	41	31	21	30	23	41	1	156398									
60	6	30	55	43	32	38	30	27	55	60	25	56	19	116116	*	10	43	27	13	31	40	33	27	25	30	18	37	13	127127									
60	1	27	38	28	27	32	28	24	28	39	24	35	43	88228	*	12	57	40	19	30	34	25	30	41	37	16	47	27	106106									
60	3	16	25	25	27	34	36	27	19	25	18	35	71	54	54	*	7	53	46	27	30	32	29	33	26	46	23	54	44	84196								
50	*	6	11	23	28	29	32	35	28	17	15	10	31	87	36	36	*	4	41	48	41	37	33	31	29	22	18	7	37	87	36	36						
50	*	6	19	25	29	26	24	22	16	13	7	24	100	22	22	*	6	31	44	47	39	34	31	22	24	27	13	45	78	46	46							
50	*	5	8	17	18	20	18	17	16	17	10	22	109	21	21	*	5	26	39	45	38	33	29	23	22	18	7	37	87	36	36							
50	*	3	6	13	12	12	16	17	14	16	10	16	114	21	21	*	5	26	35	39	34	28	25	23	20	15	8	39	91	32	32							
50	*	2	5	9	9	10	14	16	14	11	12	7	15	117	21	21	*	7	28	33	32	31	27	21	17	14	10	5	27	93	30	30						
50	*	2	4	7	7	8	11	11	9	8	9	5	9	119	21	21	*	5	27	32	30	31	29	20	12	9	6	2	31	93	29	29						
40	*	2	4	6	6	7	8	8	7	8	9	5	8	120	20	20	*	3	22	26	24	27	26	18	10	7	5	2	21	91	31	31						
40	*	2	3	7	8	8	7	6	7	8	8	4	9	122	20	20	*	3	19	22	16	20	20	15	10	7	6	3	19	91	31	31						
40	*	2	3	6	7	7	5	5	6	6	7	4	6	122	20	20	*	3	19	26	22	17	15	11	8	6	4	2	18	91	31	31						
40	*	1	2	4	5	5	5	4	4	4	5	3	5	120	20	20	*	2	19	32	29	18	12	9	7	5	1	0	19	90	32	32						
30	*	1	1	3	4	4	3	3	3	3	3	1	2	119	20	20	*	0	15	28	26	16	11	8	8	6	2	0	17	88	35	35						
30	*	1	1	3	4	4	3	3	3	3	3	1	2	119	20	20	*	1	13	24	22	13	9	8	8	7	3	1	11	88	35	35						
PER.	44	22	14.7	11	8.8	7.3	6.3	5.5	4.9	4.4	4	(DAYS)	PERIOD	44	22	14.7	11	8.8	7.3	6.3	5.5	4.9	4.4	4	(DAYS)	PER.	44	22	14.7	11	8.8	7.3	6.3	5.5	4.9	4.4	4	(DAYS)

Table A-3. Analysis of the high frequency variability of the wind at Wallops. (a) Zonal, (b) Meridional.

(a) SEASON= WINTER
POWER (M²/SEC2)

VAR N P.R.E.

SEASON= SPRING
POWER (M²/SEC2)

VÄR N P.R.E.

KM	0	60	120	180	240	300	360	420	480	540	600	660	720	780	840	900	960	1020	1080	1140	1200	1260	1320	1380	1440	1500	1560	1620	1680	1740	1800	1860	1920	1980	2040	2100	2160	2220	2280	2340	2400	2460	2520	2580	2640	2700	2760	2820	2880	2940	2980	3040	3100	3160	3220	3280	3340	3400	3460	3520	3580	3640	3700	3760	3820	3880	3940	3980	4040	4100	4160	4220	4280	4340	4400	4460	4520	4580	4640	4700	4760	4820	4880	4940	4980	5040	5100	5160	5220	5280	5340	5400	5460	5520	5580	5640	5700	5760	5820	5880	5940	5980	6040	6100	6160	6220	6280	6340	6400	6460	6520	6580	6640	6700	6760	6820	6880	6940	6980	7040	7100	7160	7220	7280	7340	7400	7460	7520	7580	7640	7700	7760	7820	7880	7940	7980	8040	8100	8160	8220	8280	8340	8400	8460	8520	8580	8640	8700	8760	8820	8880	8940	8980	9040	9100	9160	9220	9280	9340	9400	9460	9520	9580	9640	9700	9760	9820	9880	9940	9980	10040	10100	10160	10220	10280	10340	10400	10460	10520	10580	10640	10700	10760	10820	10880	10940	10980	11040	11100	11160	11220	11280	11340	11400	11460	11520	11580	11640	11700	11760	11820	11880	11940	11980	12040	12100	12160	12220	12280	12340	12400	12460	12520	12580	12640	12700	12760	12820	12880	12940	12980	13040	13100	13160	13220	13280	13340	13400	13460	13520	13580	13640	13700	13760	13820	13880	13940	13980	14040	14100	14160	14220	14280	14340	14400	14460	14520	14580	14640	14700	14760	14820	14880	14940	14980	15040	15100	15160	15220	15280	15340	15400	15460	15520	15580	15640	15700	15760	15820	15880	15940	15980	16040	16100	16160	16220	16280	16340	16400	16460	16520	16580	16640	16700	16760	16820	16880	16940	16980	17040	17100	17160	17220	17280	17340	17400	17460	17520	17580	17640	17700	17760	17820	17880	17940	17980	18040	18100	18160	18220	18280	18340	18400	18460	18520	18580	18640	18700	18760	18820	18880	18940	18980	19040	19100	19160	19220	19280	19340	19400	19460	19520	19580	19640	19700	19760	19820	19880	19940	19980	20040	20100	20160	20220	20280	20340	20400	20460	20520	20580	20640	20700	20760	20820	20880	20940	20980	21040	21100	21160	21220	21280	21340	21400	21460	21520	21580	21640	21700	21760	21820	21880	21940	21980	22040	22100	22160	22220	22280	22340	22400	22460	22520	22580	22640	22700	22760	22820	22880	22940	22980	23040	23100	23160	23220	23280	23340	23400	23460	23520	23580	23640	23700	23760	23820	23880	23940	23980	24040	24100	24160	24220	24280	24340	24400	24460	24520	24580	24640	24700	24760	24820	24880	24940	24980	25040	25100	25160	25220	25280	25340	25400	25460	25520	25580	25640	25700	25760	25820	25880	25940	25980	26040	26100	26160	26220	26280	26340	26400	26460	26520	26580	26640	26700	26760	26820	26880	26940	26980	27040	27100	27160	27220	27280	27340	27400	27460	27520	27580	27640	27700	27760	27820	27880	27940	27980	28040	28100	28160	28220	28280	28340	28400	28460	28520	28580	28640	28700	28760	28820	28880	28940	28980	29040	29100	29160	29220	29280	29340	29400	29460	29520	29580	29640	29700	29760	29820	29880	29940	29980	30040	30100	30160	30220	30280	30340	30400	30460	30520	30580	30640	30700	30760	30820	30880	30940	30980	31040	31100	31160	31220	31280	31340	31400	31460	31520	31580	31640	31700	31760	31820	31880	31940	31980	32040	32100	32160	32220	32280	32340	32400	32460	32520	32580	32640	32700	32760	32820	32880	32940	32980	33040	33100	33160	33220	33280	33340	33400	33460	33520	33580	33640	33700	33760	33820	33880	33940	33980	34040	34100	34160	34220	34280	34340	34400	34460	34520	34580	34640	34700	34760	34820	34880	34940	34980	35040	35100	35160	35220	35280	35340	35400	35460	35520	35580	35640	35700	35760	35820	35880	35940	35980	36040	36100	36160	36220	36280	36340	36400	36460	36520	36580	36640	36700	36760	36820	36880	36940	36980	37040	37100	37160	37220	37280	37340	37400	37460	37520	37580	37640	37700	37760	37820	37880	37940	37980	38040	38100	38160	38220	38280	38340	38400	38460	38520	38580	38640	38700	38760	38820	38880	38940	38980	39040	39100	39160	39220	39280	39340	39400	39460	39520	39580	39640	39700	39760	39820	39880	39940	39980	40040	40100	40160	40220	40280	40340	40400	40460	40520	40580	40640	40700	40760	40820	40880	40940	40980	41040	41100	41160	41220	41280	41340	41400	41460	41520	41580	41640	41700	41760	41820	41880	41940	41980	42040	42100	42160	42220	42280	42340	42400	42460	42520	42580	42640	42700	42760	42820	42880	42940	42980	43040	43100	43160	43220	43280	43340	43400	43460	43520	43580	43640	43700	43760	43820	43880	43940	43980	44040	44100	44160	44220	44280	44340	44400	44460	44520	44580	44640	44700	44760	44820	44880	44940	44980	45040	45100	45160	45220	45280	45340	45400	45460	45520	45580	45640	45700	45760	45820	45880	45940	45980	46040	46100	46160	46220	46280	46340	46400	46460	46520	46580	46640	46700	46760	46820	46880	46940	46980	47040	47100	47160	47220	47280	47340	47400	47460	47520	47580	47640	47700	47760	47820	47880	47940	47980	48040	48100	48160	48220	48280	48340	48400	48460	48520	48580	48640	48700	48760	48820	48880	48940	48980	49040	49100	49160	49220	49280	49340	49400	49460	49520	49580	49640	49700	49760	49820	49880	49940	49980	50040	50100	50160	50220	50280	50340	50400	50460	50520	50580	50640	50700	50760	50820	50880	50940	50980	51040	51100	51160	51220	51280	51340	51400	51460	51520	51580	51640	51700	51760	51820	51880	51940	51980	52040	52100	52160	52220	52280	52340	52400	52460	52520	52580	52640	52700	52760	52820	52880	52940	52980	53040	53100	53160	53220	53280	53340	53400	53460	53520	53580	53640	53700	53760	53820	53880	53940	53980	54040	54100	54160	54220	54280	54340	54400	54460	54520	54580	54640	54700	54760	54820	54880	54940	54980	55040	55100	55160	55220	55280	55340	55400	55460	55520	55580	55640	55700	55760	55820	55880	55940	55980	56040	56100	56160	56220	56280	56340	56400	56460	56520	56580	56640	56700	56760	56820	56880	56940	56980	57040	57100	57160	57220	57280	57340	57400	57460	57520	57580	57640	57700	57760	57820	57880	57940	57980	58040	58100	58160	58220	58280	58340	58400	58460	58520	58580	58640	58700	58760	58820	58880	58940	58980	59040	59100	59160	59220	59280	59340	59400	59460	59520	59580	59640	59700	59760	59820	59880	59940	59980	60040	60100	60160	60220	60280	60340	60400	60460	60520	60580	60640	60700	60760	60820	60880	60940	60980	61040	61100	61160	61220	61280	61340	61400	61460	61520	61580	61640	61700	61760	61820	61880	61940	61980	62040	62100	62160	62220	62280	62340	62400	62460	62520	62580	62640	62700	62760	62820	62880	62940	62980	63040	63100	63160	63220	63280	63340	63400	63460	63520	63580	63640	63700	63760	63820	63880	63940	63980	64040	64100	64160	64220	64280	64340	64400	64460	64520	64580	64640	64700	64760	64820	64880	64940	64980	65040	65100	65160	65220	65280	65340	65400	65460	65520	65580	65640	65700	65760	65820	65880	65940	65980	66040	66100	66160	66220	6628

(a) SEASON= WINTER POWER (M2/SEC2)												SEASON= SPRING POWER (M2/SEC2)																				
	VAR	N	P.R.E.		VAR	N	P.R.E.		VAR	N	P.R.E.																					
KM *				*				*																								
60	102	369	464	348	251	289	302	171	99	115	60	322	90	31	31	*	26	98	117	99	80	58	57	97	126	110	51	150	140	19	92	
	92	368	479	366	263	295	308	183	113	117	57	407	126	20	20	*	29	88	106	92	69	50	50	72	98	103	54	109	180	18	18	
	97	372	482	383	273	283	285	178	132	131	57	359	154	19	19	*	33	74	86	78	57	46	45	42	58	86	53	91	209	18	18	
	118	384	468	373	267	262	254	152	120	132	60	385	171	19	19	*	35	68	73	70	61	53	46	35	39	63	42	83	233	17	17	
	127	399	465	393	252	246	233	139	99	113	59	335	123	18	18	*	34	64	68	69	67	56	41	29	31	49	33	77	251	17	17	
	141	423	477	345	239	217	192	128	101	107	56	349	190	18	18	*	31	61	68	67	65	52	36	27	30	49	33	73	260	17	17	
	156	431	475	344	235	195	158	118	110	111	59	340	195	18	18	*	31	63	68	61	54	46	39	31	32	47	29	73	268	17	17	
	163	404	428	320	227	178	142	115	105	107	61	329	198	18	18	*	30	64	70	63	53	47	44	36	34	36	18	67	271	17	17	
	170	380	382	283	201	159	131	103	84	82	48	283	202	18	18	*	31	65	74	69	59	49	41	31	28	26	11	76	273	17	17	
	173	365	356	240	164	141	117	87	67	55	29	263	205	18	18	*	33	63	68	59	51	42	32	24	22	20	10	58	274	16	16	
	40	159	339	326	205	134	118	99	81	65	40	16	224	209	18	18	*	33	59	58	45	38	32	26	21	20	18	9	52	276	16	16
	139	293	290	185	117	94	74	68	62	36	11	199	213	18	18	*	30	57	55	40	33	29	24	19	19	17	8	46	281	16	16	
	120	245	250	166	100	71	51	49	50	35	11	167	215	18	18	*	26	52	52	39	31	27	22	18	18	16	7	45	284	16	16	
	98	192	199	135	78	52	38	34	34	27	10	130	213	18	18	*	24	41	42	36	30	24	19	17	17	15	7	38	286	16	16	
	74	142	145	99	56	35	28	27	25	20	8	94	212	18	18	*	16	32	34	31	26	19	15	14	13	12	6	31	287	16	16	
	30	65	117	119	83	46	26	22	25	23	17	6	65	214	18	18	*	15	31	33	29	23	16	13	12	11	10	5	26	287	16	16
*****																									*****							
PER.	44	22	14.7	11	8.8	7.3	6.3	5.5	4.9	4.4	4	(DAYS)					PERIOD	44	22	14.7	11	8.8	7.3	6.3	5.5	4.9	4.4	4	(DAYS)			

SEASON= SUMMER POWER (MH2/SEC2)										SEASON= AUTUMN POWER (MH2/SEC2)										SEASON= WINTER POWER (MH2/SEC2)													
VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.																									
KM	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•				
0	2	65	112	122	104	95	93	94	107	85	31	150	137	19	19	•	•	14	114	155	157	171	149	114	116	97	78	45	168	127	19	19	
0	6	66	102	102	92	89	87	91	99	75	28	105	188	17	17	•	•	15	106	149	153	160	138	107	112	100	82	45	170	180	18	18	
0	9	57	82	74	70	79	80	85	85	68	31	110	229	17	17	•	•	20	104	152	156	146	117	94	107	108	90	45	146	213	18	18	
0	7	38	59	56	54	66	74	75	69	64	35	81	253	17	17	•	•	30	117	168	164	140	109	91	107	112	96	47	181	227	17	17	
0	5	27	40	42	42	55	66	64	53	50	28	65	268	16	16	•	•	35	120	172	152	114	98	93	97	93	85	45	162	238	17	17	
50	•	3	22	34	38	39	45	54	52	41	38	22	55	279	16	16	•	•	38	110	150	119	75	77	84	75	61	32	123	252	17	17	
0	2	19	32	37	37	36	40	34	32	29	16	43	286	16	16	•	•	43	98	112	86	58	64	67	57	45	39	22	92	264	17	17	
0	3	18	29	33	33	30	31	32	24	24	12	37	288	16	16	•	•	44	89	91	72	53	55	55	48	42	37	19	87	267	16	16	
0	2	16	25	28	28	25	24	29	29	24	11	35	290	16	16	•	•	47	87	85	69	54	51	49	43	37	31	16	81	271	16	16	
0	3	17	25	28	29	24	19	23	26	22	10	28	294	16	16	•	•	48	82	78	68	58	49	44	39	31	24	12	75	275	16	16	
40	•	3	17	24	28	31	26	19	19	20	18	9	35	300	16	16	•	•	46	84	79	68	52	37	34	33	29	25	13	72	278	16	16
0	2	12	18	22	24	20	17	15	15	14	7	21	303	16	16	•	•	48	92	84	63	40	25	22	25	28	27	14	69	277	16	16	
0	1	7	13	17	15	13	14	13	11	11	6	15	304	16	16	•	•	44	44	76	54	33	20	16	20	26	24	11	59	277	16	16	
0	0	6	13	16	13	11	13	12	11	10	5	17	303	16	16	•	•	36	45	58	42	29	20	17	19	21	9	47	277	16	16		
0	1	6	12	14	12	10	10	9	9	8	4	12	302	16	16	•	•	29	51	43	33	25	18	17	17	17	10	39	277	16	16		
30	•	1	6	11	12	10	8	7	7	7	3	10	304	16	16	•	•	27	46	38	29	22	16	15	15	15	17	10	33	278	16	16	
PERIOD	44	22	14.7	11	8.8	7.0	6.3	5.5	4.9	4.4	4	(DAYS)	PERIOD	44	22	14.7	11	8.8	7.3	6.3	5.5	4.9	4.4	4	(DAYS)								

(b) SEASON= WINTER POWER (M2/SEC2)										SEASON= SPRING POWER (M2/SEC2)									
	VAR	N	P.R.E.		VAR	N	P.R.E.		VAR	N	P.R.E.								
KM *					*				*										
60	1 107 158 162 156 132 115 117 115 114 61 168 90 31 31	*	2 31 58 66 63 70 77 78 75 63 30 95 140 19 19																
*	11 118 170 167 152 122 98 101 107 99 49 176 126 20 20	*	1 29 51 57 55 63 70 71 70 66 34 76 180 18 18																
*	29 131 167 151 130 100 71 81 100 90 44 146 150 19 19	*	0 26 42 43 44 51 57 59 62 33 68 209 18 18																
*	43 138 148 125 107 79 61 76 95 98 53 146 165 19 19	*	2 23 36 36 38 44 45 45 46 44 23 52 233 17 17																
*	44 137 146 129 108 74 65 74 82 85 47 138 177 19 19	*	6 22 29 30 33 40 40 37 37 38 21 41 251 17 17																
50	42 133 151 138 110 77 71 76 77 68 32 136 185 18 18	*	6 21 27 29 33 40 38 34 36 42 25 91 260 17 17																
*	44 129 148 135 106 77 71 73 74 59 24 138 189 18 18	*	6 22 30 33 37 40 35 29 32 39 23 44 268 17 17																
*	37 109 130 122 96 73 68 64 60 48 20 117 192 18 18	*	6 22 32 36 36 36 32 27 29 33 19 45 271 17 17																
*	28 90 111 99 77 61 60 56 48 41 21 194 193 18 18	*	5 21 32 34 32 30 29 27 27 26 14 39 273 17 17																
*	25 78 90 74 61 50 49 48 42 39 21 64 195 18 18	*	5 20 29 28 26 27 26 24 20 17 8 32 274 16 16																
40	21 58 66 56 53 43 35 33 33 34 19 62 200 18 18	*	6 19 25 23 22 23 23 19 15 12 6 26 276 16 16																
*	14 40 50 51 48 36 25 22 26 28 14 46 205 18 18	*	6 16 19 20 20 20 19 15 12 12 6 23 281 16 16																
*	8 30 44 47 40 32 23 21 23 23 12 45 209 18 18	*	6 13 15 16 16 16 13 12 13 12 7 19 284 16 16																
*	5 22 33 35 30 26 22 20 19 20 11 34 208 18 18	*	5 12 13 13 12 12 14 13 12 12 7 18 286 16 16																
*	5 16 23 23 21 20 17 13 12 15 10 23 205 18 18	*	3 10 12 11 9 10 12 10 9 10 6 13 287 16 16																
30	6 15 19 18 16 17 15 10 8 12 8 17 208 18 18	*	2 8 11 10 9 10 10 8 6 8 5 11 287 16 16																
*																			
PERIOD	4.6 22 14.7 11 8.8 7.3 6.3 5.5 4.9 4.4 4 (DAYS)				PERIOD	4.6 22 14.7 11 8.8 7.3 6.3 5.5 4.9 4.4 4 (DAYS)													

SEASON= SUMMER POWER (M2/SEC2)												SEASON= AUTUMN POWER (M2/SEC2)																						
KM	*												*																					
60	*	1	41	65	76	76	69	61	52	61	67	33	88	137	19	19	*	3	48	79	92	94	84	78	81	72	77	49	115	127	19	19		
*	*	1	39	65	72	66	58	55	50	56	62	31	79	188	17	17	*	1	44	71	88	91	80	73	73	69	74	44	94	179	18	18		
*	*	0	33	59	65	55	48	51	48	46	50	26	63	229	17	17	*	8	45	62	77	83	74	63	57	59	61	33	92	213	18	18		
*	*	0	26	47	54	53	52	50	44	39	38	20	64	253	17	17	*	11	51	65	63	66	65	52	44	44	41	21	70	226	17	17		
50	*	1	22	37	39	43	47	42	38	36	32	16	46	268	16	16	*	12	53	65	56	55	54	43	36	30	13	60	234	17	17			
*	*	2	21	31	29	30	36	35	33	33	28	13	41	279	16	16	*	9	49	64	58	56	50	38	33	35	29	11	62	249	17	17		
*	*	3	19	30	27	25	29	30	28	27	24	11	36	286	16	16	*	7	44	61	57	53	49	37	30	31	26	11	57	261	17	17		
*	*	2	16	26	26	24	23	22	22	22	20	10	30	288	16	16	*	8	40	52	48	47	44	34	29	29	25	11	50	264	16	16		
*	*	2	12	18	22	22	18	17	19	19	17	8	24	290	16	16	*	6	34	44	42	43	38	28	26	26	24	13	47	271	16	16		
40	*	2	8	13	16	17	15	14	16	18	16	7	19	294	16	16	*	4	32	41	37	38	32	21	20	18	17	10	35	275	15	67		
*	*	1	7	11	13	14	13	12	12	14	14	7	16	299	16	16	*	5	29	36	31	31	25	15	15	14	11	6	31	278	16	16		
*	*	1	7	10	11	12	12	11	9	10	11	5	13	303	16	16	*	5	23	26	23	23	18	12	14	15	12	6	23	277	16	16		
*	*	1	6	9	9	9	10	9	9	9	9	5	12	304	16	16	*	5	18	22	19	18	13	12	15	15	11	5	21	276	16	16		
*	*	0	4	7	8	8	8	8	8	8	8	4	9	303	16	16	*	3	15	20	18	16	12	11	14	13	9	4	19	277	16	16		
30	*	0	3	6	7	7	7	7	7	7	6	3	8	302	16	16	*	2	11	15	16	16	12	10	11	10	9	5	16	276	16	16		
*	*	*	*	*	*	*	*	*	*	*	*	*	*	*	*	*	*	1	9	13	14	15	12	9	9	9	9	5	12	277	16	16		
*****												*****																						
PFR	44	22	14	7	11	8	8	7	3	6	3	5	4	4	4	4	(DAYS)	PERIOD	44	22	14	7	11	8	8	7	3	6	3	5	4	4	4	(DAYS)

Table A-5. Analysis of the high frequency variability of the wind at White Sands. (a) Zonal, (b) Meridional.

(a)

SEASON= WINTER POWER (M ² /SEC2)												SEASON= SPRING POWER (M ² /SEC2)												SEASON= SUMMER POWER (M ² /SEC2)							
KH	VAR	N	P.R.F.	VAR	N	P.R.F.	VAR	N	P.R.F.	VAR	N	P.R.F.	VAR	N	P.R.F.	VAR	N	P.R.F.	VAR	N	P.R.F.	VAR	N	P.R.F.							
60	181	354	278	236	205	118	100	119	84	7	20	243	12	120120	14	78	111	89	60	50	48	50	60	76	45	.98	8	124124			
60	171	336	268	210	185	112	92	104	87	29	3	231	58	66	16	77	100	77	61	57	48	45	54	65	37	.93	62	62	62		
60	161	308	246	175	163	111	87	96	85	56	20	207	107	21	21	71	78	57	55	60	45	34	42	47	26	74	115	21	21		
50	155	297	241	177	165	116	93	102	85	65	31	223	147	19	19	68	66	46	45	48	35	25	33	33	15	.58	165	19	19		
50	146	286	257	197	173	113	89	102	86	67	32	226	179	19	19	32	71	68	49	41	35	24	22	29	25	10	.59	191	19	19	
50	139	281	277	226	188	111	81	92	79	53	21	216	200	18	18	35	72	72	52	38	28	19	20	25	24	11	.56	204	19	19	
50	128	265	290	266	207	108	80	92	73	40	12	231	215	18	18	37	71	70	52	35	23	18	20	22	24	14	.55	214	18	18	
50	114	242	281	269	200	96	80	95	70	35	10	219	227	18	18	36	65	55	52	37	25	21	21	20	20	12	.54	220	18	18	
50	111	231	256	236	173	84	71	81	60	29	6	189	231	18	18	30	55	59	55	44	29	22	21	17	16	9	.51	227	18	18	
50	120	231	228	192	141	75	62	59	43	21	2	143	231	18	18	22	44	55	59	47	27	18	18	16	14	7	.47	232	18	18	
40	126	224	197	147	105	61	51	41	27	16	4	143	232	18	18	17	40	54	56	42	22	13	15	16	14	7	.42	234	18	18	
40	114	207	168	107	74	47	39	26	16	8	8	120	233	18	18	17	41	52	50	36	20	13	13	14	14	7	.38	232	18	18	
40	97	177	137	77	56	41	31	19	12	15	9	95	233	18	18	20	42	47	44	35	21	14	13	12	13	8	.40	229	18	18	
40	74	135	101	57	45	37	28	18	14	15	8	79	230	18	18	20	36	45	34	30	21	14	12	11	12	7	.35	229	18	18	
40	54	93	69	42	37	33	25	17	14	14	7	56	225	18	18	15	25	42	44	23	17	11	9	10	11	6	.23	226	18	18	
30	43	74	56	36	34	31	23	16	13	13	7	43	221	18	18	11	20	21	21	19	15	10	8	9	9	5	.19	224	18	18	
PER.	44	22	14.7	11	8.8	7.3	6.3	5.5	4.9	4.4	4	(DAYS)				PERIOD	44	22	14.7	11	8.8	7.3	6.3	5.5	4.9	4.4	4	(DAYS)			

SEASON= SUMMER POWER (M ² /SEC2)												SEASON= AUTUMN POWER (M ² /SEC2)												SEASON= WINTER POWER (M ² /SEC2)							
KH	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.							
60	23	32	86	112	99	97	122	119	87	76	44	115	18	109204	34	115	97	73	67	35	34	102	122	63	17	122	24	109244			
60	11	45	89	105	94	97	120	113	79	63	35	120	68	54	54	41	120	101	68	63	42	34	78	92	55	21	89	58	66156		
60	5	50	93	100	78	77	101	101	76	55	25	106	136	20	20	44	116	102	66	62	51	41	58	62	48	24	100	94	28	28	
60	6	51	91	96	67	56	76	84	75	70	37	93	179	19	19	40	101	96	70	61	48	45	58	59	49	25	92	125	20	20	
50	13	53	74	74	60	52	59	61	59	74	45	106	199	19	19	41	91	92	68	53	42	39	46	47	64	24	86	156	20	20	
50	10	33	44	47	44	40	42	44	43	48	28	45	213	18	18	42	85	87	64	49	41	33	33	31	30	18	71	177	19	19	
50	4	17	27	33	30	27	30	35	35	31	15	41	224	18	18	42	76	78	62	46	35	27	27	26	21	11	67	184	19	19	
50	3	14	23	26	24	24	27	28	27	26	14	33	231	18	18	30	66	70	57	37	27	24	24	24	9	55	191	19	19		
50	3	15	23	25	25	26	27	24	21	20	11	26	235	18	18	3	62	69	53	31	24	25	22	20	19	10	54	200	19	19	
40	4	18	26	27	29	30	30	26	21	17	8	39	237	18	18	34	62	72	54	30	27	26	18	14	16	9	55	206	18	18	
40	4	16	23	24	25	27	28	24	19	17	9	29	238	18	18	33	57	67	56	32	25	23	18	14	12	6	46	211	18	18	
40	2	12	19	20	20	20	24	18	16	18	10	23	238	18	18	32	51	59	55	34	22	17	17	15	10	4	48	214	18	18	
40	2	9	15	19	20	17	15	14	13	15	9	20	238	18	18	28	51	49	47	31	17	12	14	13	10	5	39	213	19	19	
40	1	7	12	18	20	16	14	12	12	7	3	18	237	18	18	19	36	40	35	24	13	10	12	12	7	31	213	19	19		
30	1	6	10	14	14	13	12	13	7	6	3	18	218	18	18	11	26	32	27	21	18	11	11	12	7	24	212	19	19		
PER.	44	22	14.7	11	8.8	7.3	6.3	5.5	4.9	4.4	4	(DAYS)				PERIOD	44	22	14.7	11	8.8	7.3	6.3	5.5	4.9	4.4	4	(DAYS)			

SEASON= SUMMER POWER (M ² /SEC2)												SEASON= AUTUMN POWER (M ² /SEC2)												SEASON= WINTER POWER (M ² /SEC2)						
KH	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.	VAR	N	P.R.E.															
60	3	43	65	75	73	74	81	80	74	66	33	108	18	112112	16	46	36	47	69	58	36	34	32	33	21	77	24	109109		
60	1	42	58	63	63	65	70	69	64	60	31	75	68	54	54	12	38	35	45	55	46	37	35	31	30	18	42	58	67	67
60	5	37	50	51	53	54	50	48	49	50	26	67	136	20	20	8	33	38	40	35	32	41	44	36	30	15	51	94	28	28
60	4	29	43	47	49	47	41	41	37	17	55	177	19	19	8	38	43	36	28	31	43	45	40	33	16	51	125	20	20	
50	2	21	34	40	41	39	38	40	39	29	13	45	196	19	19	8	37	43	36	32										

(a) SEASON= WINTER POWER (M2/SEC2)												SEASON= SPRING POWER (M2/SEC2)											
	VAR	N	P.R.E.		VAR	N	P.R.E.																
KM																							
60	14 251 364 280 272 254 167 129 83 102 84 282 7 139311	15 66 85 62 49 49 49 65 66 40 13 79 7 139139																					
	8 261 335 266 279 240 155 139 103 119 91 260 21 114280	18 68 82 61 52 49 41 53 59 41 15 78 25 103103																					
	44 261 295 271 304 235 147 123 94 134 103 299 60 86 66 66	22 61 69 54 54 47 36 42 48 37 15 68 80 44 44																					
	64 236 255 268 309 241 150 82 47 110 91 261 85 39 39	21 49 56 52 53 42 31 38 41 31 12 59 119 22 22																					
	59 197 205 225 277 235 156 74 39 89 69 218 100 23 23	18 45 58 58 52 36 23 27 33 30 14 56 140 22 22																					
50	45 171 183 197 243 233 180 46 65 100 62 .217 111 23 23	20 48 66 65 51 31 17 20 26 29 16 54 146 22 22																					
	57 190 197 196 220 219 182 101 74 104 60 231 115 23 23	23 50 68 71 51 29 17 18 21 27 17 58 150 22 22																					
	93 239 224 196 194 172 145 88 63 89 53 218 116 23 23	23 43 62 71 50 30 23 20 18 23 15 54 152 21 21																					
	122 272 238 178 166 113 106 83 62 73 43 211 117 22 22	19 32 49 62 48 33 26 21 16 17 11 50 156 21 21																					
	123 266 229 141 93 73 79 76 68 67 35 177 118 22 22	16 26 37 50 44 32 24 19 13 11 7 37 160 21 21																					
40	112 233 196 102 56 49 58 68 71 61 27 149 119 22 22	14 24 33 43 40 30 22 16 10 8 5 34 162 21 21																					
	94 188 152 76 43 36 40 55 66 53 21 118 117 23 23	11 23 33 41 36 27 21 15 9 8 5 33 161 21 21																					
	70 139 113 65 42 29 28 39 51 43 16 88 116 23 23	11 23 32 38 31 21 18 13 8 7 4 29 161 21 21																					
	51 107 94 63 43 27 21 27 37 31 11 70 116 23 23	11 23 29 34 26 16 14 11 7 5 2 25 161 21 21																					
	43 94 84 56 42 28 15 16 26 21 6 62 116 23 23	11 22 26 29 23 13 11 9 6 4 2 22 160 21 21																					
30	40 89 78 51 40 30 13 10 20 17 5 49 116 23 23	9 21 25 26 21 11 9 7 5 4 2 17 160 21 21																					

PER.	44 22 14.7 11 8.8 7.3 6.3 5.5 4.9 4.4 4 (DAYS)	PERIOD 44 22 14.7 11 8.8 7.3 6.3 5.5 4.9 4.4 4 (DAYS)																					

SEASON= SUMMER POWER (M2/SEC2)										SEASON= AUTUMN POWER (M2/SEC2)										SEASON= WINTER POWER (M2/SEC2)											
	VAR	N	P.R.E.		VAR	N	P.R.E.		VAR	N	P.R.E.		VAR	N	P.R.E.		VAR	N	P.R.E.		VAR	N	P.R.E.		VAR	N	P.R.E.		VAR	N	P.R.E.
KH =										*											*										
60 =	35	40	83	100	165	179	126	93	81	71	36	127	15	129324	*	31	143	157	93	58	64	69	65	73	69	32	117	11	154513		
*	25	39	74	86	129	154	119	85	76	73	38	130	35	91	91	*	29	145	167	101	59	57	57	60	73	65	27	121	47	79	79
*	8	33	62	62	70	90	82	63	63	63	33	81	96	26	26	*	18	128	171	116	73	57	39	44	64	53	18	108	87	36	36
*	1	23	44	50	43	39	40	*1	43	44	26	50	136	21	21	*	12	106	161	126	86	50	27	27	48	38	11	97	118	22	22
*	3	17	33	42	36	28	28	27	25	29	19	40	159	17	62	*	18	100	156	126	78	50	21	18	35	27	7	9	135	21	21
50 =	4	14	24	28	29	29	27	24	20	20	12	32	168	19	19	*	21	97	144	111	60	35	17	14	24	23	10	80	150	20	20
*	4	13	18	20	23	25	24	21	18	17	9	27	187	19	19	*	20	84	116	82	41	24	16	14	19	24	14	64	159	20	20
*	4	12	16	18	19	20	19	18	16	13	7	21	190	19	19	*	17	65	84	59	29	18	17	16	18	22	13	47	161	20	20
*	3	10	15	17	18	18	19	16	12	6	21	193	19	19	*	14	50	61	44	24	17	18	16	16	20	12	41	163	20	20	
*	2	10	15	17	17	16	18	19	15	13	7	21	194	19	19	*	13	37	42	33	21	16	15	15	19	12	33	166	20	20	
40 =	3	10	15	17	16	14	15	15	12	11	7	20	195	19	19	*	11	28	31	26	17	12	12	13	14	17	10	25	165	20	20
*	7	8	13	14	13	12	11	9	8	9	5	14	194	19	19	*	9	24	29	24	17	12	11	12	12	13	7	24	166	20	20
*	1	6	10	11	10	9	8	7	6	7	5	10	189	20	20	*	7	23	29	21	15	13	12	11	10	5	22	166	20	20	
*	0	5	9	10	10	8	7	6	6	7	4	10	181	20	20	*	5	21	27	18	13	13	12	11	8	8	5	20	166	20	20
*	0	4	8	10	10	8	7	6	5	5	3	9	173	20	20	*	4	16	21	14	11	11	11	10	8	8	5	16	165	19	87
30 =	0	4	8	10	9	8	8	6	5	5	2	8	170	20	20	*	3	13	18	13	10	10	9	9	6	9	5	13	165	20	20
PER. =	44	22	14.7	11	8.8	7.3	6.3	5.5	4.9	4.4	4	(DAYS)	*****														*****				
PERIOD	44	22	14.7	11	8.8	7.3	6.3	5.5	4.9	4.4	4	(DAYS)	*****														*****				

Table A-7. Analysis of the high frequency variability of the wind at Hawaii. (a) Zonal, (b) Meridional.

(a) SEASON= WINTER
POWER (M2/SEC2)

VAR N P.R.E.

KM	0	30	79	92	84	77	64	56	73	72	62	33	95	13	122127
60	0	31	78	89	77	71	62	55	91	84	75	44	110	58	66150
50	0	31	76	94	74	64	65	83	111	95	85	52	121	114	20 89
40	0	32	73	100	81	59	70	101	115	90	72	39	122	149	22 22
30	0	34	65	89	82	59	71	107	111	77	55	28	112	160	21 21
20	0	31	58	78	78	64	66	86	90	71	53	25	102	166	21 21
10	0	27	59	76	70	56	51	56	67	72	63	29	86	167	21 21
0	0	20	58	79	66	46	41	43	52	66	71	37	85	168	21 21
40	0	7	32	50	46	41	47	41	59	48	60	32	73	169	21 21
30	0	8	31	41	36	38	45	37	26	25	28	15	52	149	22 22
20	0	7	27	37	32	31	34	30	23	22	23	13	35	171	21 21
10	0	5	14	26	25	25	26	26	25	21	19	11	34	171	21 21
0	0	2	13	20	21	22	21	22	23	19	15	8	23	170	21 21
30	0	0	11	21	21	19	17	17	17	14	12	6	21	170	21 21
20	0	0	11	21	21	18	16	16	14	13	12	6	19	170	21 21
10	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0

SEASON= SPRING
POWER (M2/SEC2)

VAR N P.R.E.

60	0	24	104	133	63	15	16	15	15	15	14	21	16	59	20 116253
50	0	17	97	134	68	16	17	20	21	15	22	17	17	63	52 79231
40	0	11	92	134	73	21	20	28	27	16	20	16	64	116 20 89	
30	0	16	97	130	71	27	23	28	26	15	15	11	64	145 22111	
20	0	22	98	121	68	31	23	25	24	17	13	7	63	162 20 89	
10	0	22	89	112	69	33	24	24	22	16	13	8	60	167 21 21	
0	0	17	75	105	73	34	23	24	20	14	15	10	58	168 21 21	
40	0	15	63	92	68	33	22	22	21	17	18	10	54	168 21 21	
30	0	14	52	74	54	27	21	22	22	19	17	9	50	170 21 21	
20	0	9	40	50	40	21	18	18	18	16	8	33	170	19 85	
10	0	6	29	44	36	23	17	17	19	17	16	8	33	169 21 21	
0	0	5	23	41	44	29	16	15	16	15	7	31	170	21 21	
40	0	7	27	37	32	31	23	22	23	13	15	12	13	6	33 170 21 21
30	0	5	14	26	25	26	20	19	19	18	18	16	12	6	24 170 21 21
20	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
10	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0

PER. 44 22 14.7 11 8.8 7.3 6.3 5.5 4.9 4.4 4 (DAYS)

PERIOD 44 22 14.7 11 8.8 7.3 6.3 5.5 4.9 4.4 4 (DAYS)

SEASON= SUMMER
POWER (M2/SEC2)

VAR N P.R.E.

KM	0	29	91	111	115	92	54	35	31	46	51	22	86	15	119119
60	0	32	94	114	112	89	55	37	35	50	53	24	104	60	64 64
50	0	39	91	104	98	82	58	40	42	54	52	24	100	124	22 22
40	0	41	77	88	90	77	57	42	44	53	48	22	87	150	22 22
30	0	39	72	92	102	81	54	38	37	47	49	25	92	159	21 21
20	0	38	77	102	107	86	56	33	27	37	48	28	92	165	21 21
10	0	37	78	98	96	84	61	36	27	33	42	25	89	170	21 21
0	0	34	75	88	79	76	61	38	30	33	39	23	85	172	21 21
40	0	33	74	81	67	67	55	33	26	27	33	20	72	173	21 21
30	0	30	68	72	61	65	52	29	24	22	22	13	67	176	21 21
20	0	25	55	59	56	59	46	31	27	21	17	8	56	177	21 21
10	0	22	44	54	56	49	34	26	23	20	19	9	57	178	21 21
0	0	19	41	52	54	40	22	14	15	18	19	9	45	180	21 21
40	0	14	34	46	46	32	15	10	11	14	15	8	31	181	21 21
30	0	9	25	37	40	26	12	11	12	10	11	6	32	181	21 21
20	0	6	20	32	36	24	11	12	12	9	10	6	17	181	21 21
10	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0

SEASON= AUTUMN
POWER (M2/SEC2)

VAR N P.R.E.

60	0	4	36	60	59	61	60	48	37	41	67	47	47	69	8 129129	
50	0	8	53	78	68	58	58	49	34	38	66	46	74	48	77 77	
40	0	27	90	116	87	63	61	51	34	38	60	39	97	110	22 22	
30	0	36	112	137	93	67	64	49	33	36	46	27	111	150	21 21	
20	0	36	109	130	85	60	55	40	30	32	33	18	80	173	20 20	
10	0	34	100	116	79	58	49	37	34	34	31	15	88	184	20 20	
0	0	29	87	99	69	57	47	35	33	35	32	16	80	189	20 20	
40	0	34	75	88	79	76	61	38	30	33	39	23	13	80	189	20 20
30	0	33	74	81	67	67	55	33	26	27	33	22	23	13	84 190 20 20	
20	0	30	68	72	61	65	52	29	24	22	22	13	13	23	13 190 20 20	
10	0	25	55	59	56	59	46	31	27	22	22	13	13	23	13 190 20 20	
0	0	21	46	54	56	49	34	26	23	22	22	13	13	23	13 190 20 20	
40	0	18	41	52	54	40	22	14	15	18	19	9	45	180	21 21	
30	0	17	39	52	56	49	34	26	23	22	22	13	13	23	13 190 20 20	
20	0	14	32	45	47	40	24	17	18	21	21	11	11	23	13 190 20 20	
10	0	13	29	42	44	37	24	17	18	21	21	11	11	23	13 190 20 20	
0	0	10	24	34	37	29	24	17	18	21	21	11	11	23	13 190 20 20	
40	0	9	23	37	40	33	27	21	21	21	21	11	11	23	13 190 20 20	
30	0	8	21	34	37	30	24	17	18	21	21	11	11	23	13 190 20 20	
20	0	5	16	21	24	21	17	18	21	21	21	11	11	23	13 190 20 20	
10	0	4	13	19	21	17	18	21	21	21	21	11	11	23	13 190 20 20	
0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	

PER. 44 22 14.7 11 8.8 7.3 6.3 5.5 4.9 4.4 4 (DAYS)

PERIOD 44 22 14.7 11 8.8 7.3 6.3 5.5 4.9 4.4 4 (DAYS)

SEASON= SUMMER
POWER (M2/SEC2)

VAR N P.R.E.

KM	0	3	61	85	68	65	83	75	48	36	38	22	92	15	119119
60	0	4	48	72	68	61	65	60	43	33	32	17			